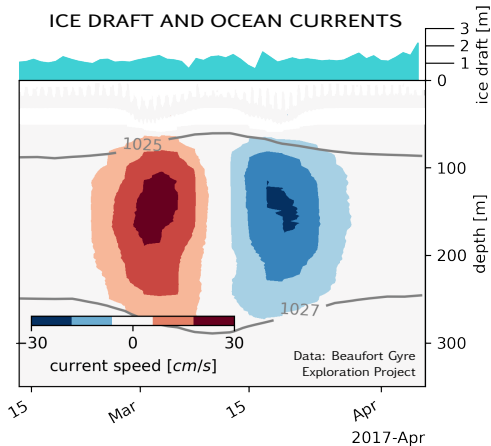


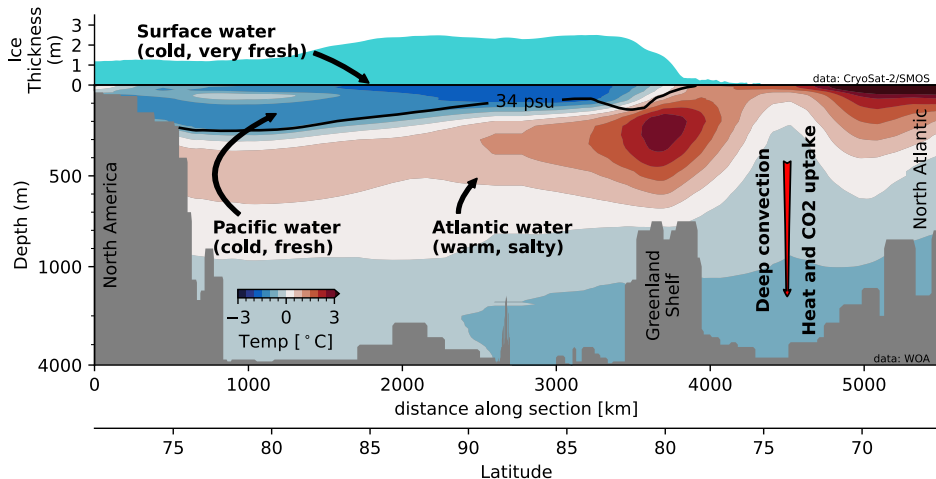
WHAT CONTROLS  
MESOSCALE EDDIES  
AND EDDY FLUXES  
IN THE ARCTIC OCEAN?

GIANLUCA MENEGHELLO



# TO EACH WATER MASS THEIR OWN EDDY SYSTEM

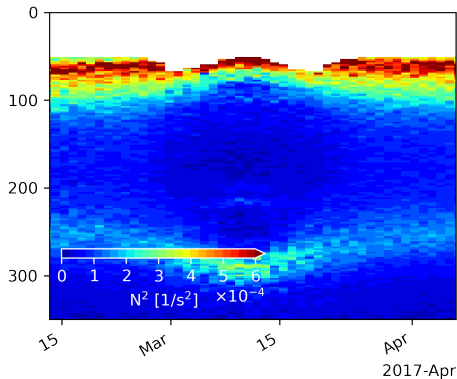
Independent, very different mesoscale eddies develop in each water mass.



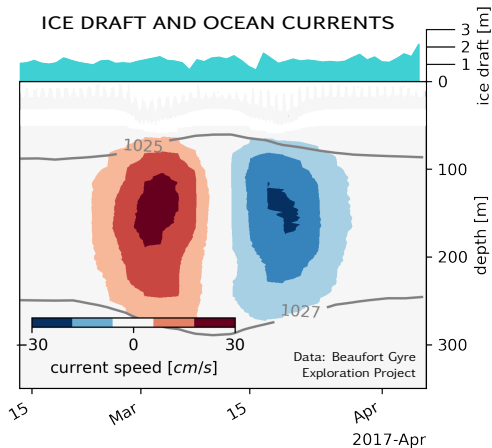
# OBSERVATIONS OF MESOSCALE EDDIES IN THE CANADIAN ARCTIC

Peaks in stratification limit vertical extent of mesoscale eddies

STRATIFICATION

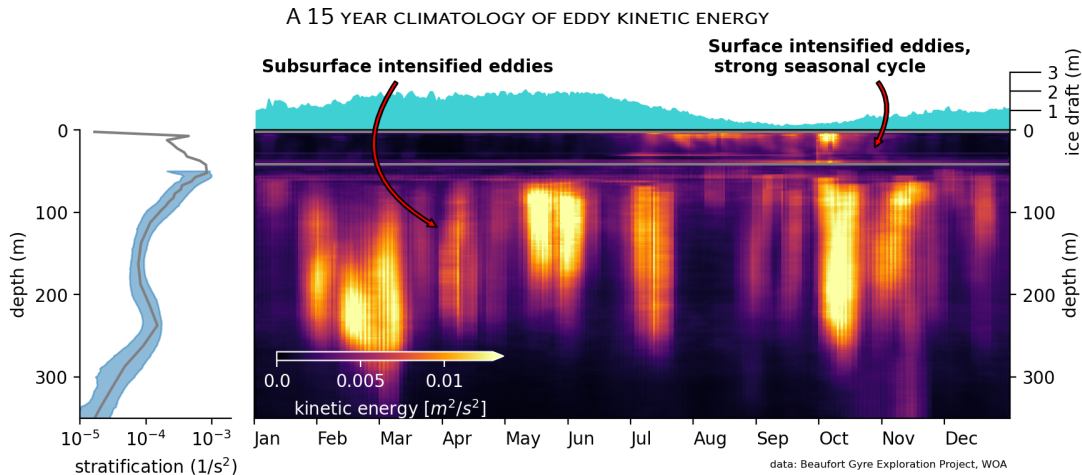


ICE DRAFT AND OCEAN CURRENTS



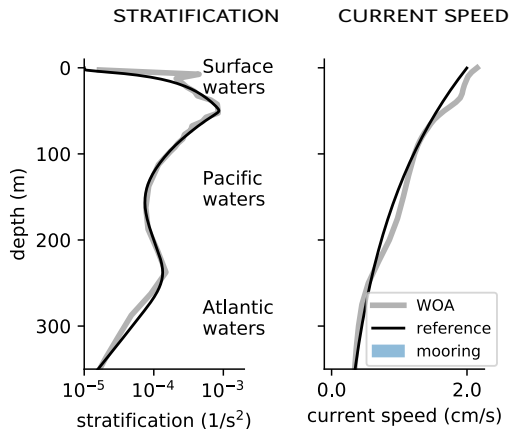
# OBSERVATIONS OF MESOSCALE EDDIES IN THE CANADIAN ARCTIC

Sea ice controls the level of kinetic energy at the surface



# THE ORIGIN OF MESOSCALE EDDIES

Can baroclinic instability explain the vertical structure of the Arctic's eddy field?



## Quasi-geostrophic balance

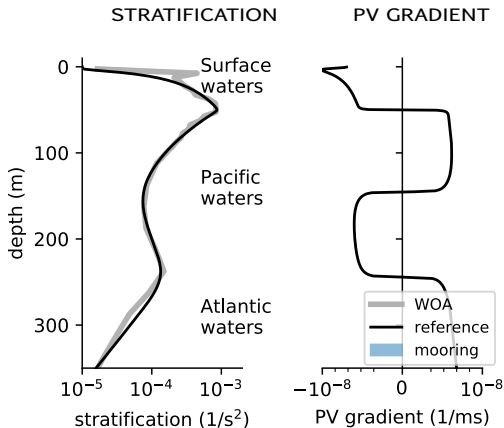
$$\frac{d\rho}{dt} = -\mathbf{u} \cdot \nabla \bar{\rho} \pm w_{Ek} \frac{\partial \bar{\rho}}{\partial z} \quad \text{Density conservation (boundary)}$$

$$\frac{dq}{dt} = -\mathbf{u} \cdot \nabla Q \quad \text{PV conservation (interior)}$$

Ekman pumping  $w_{Ek}$  has a stabilizing effect on baroclinic instability.

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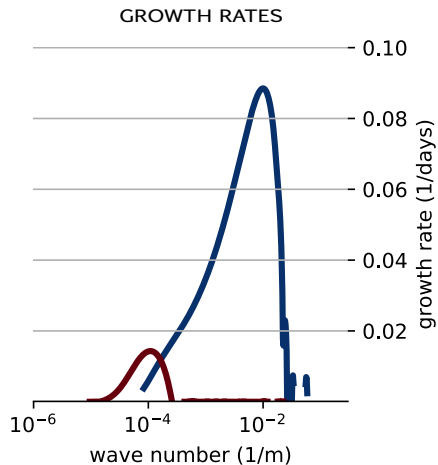
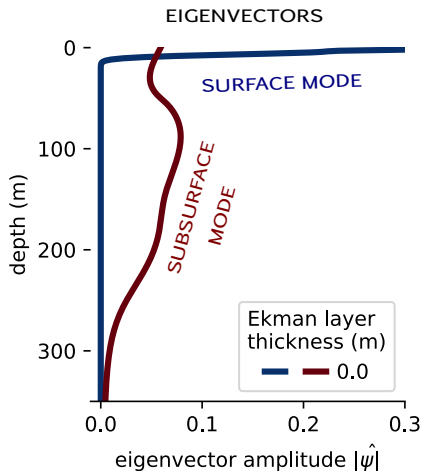
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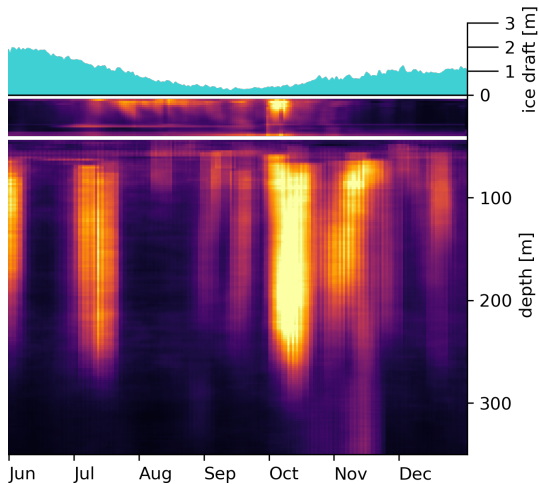
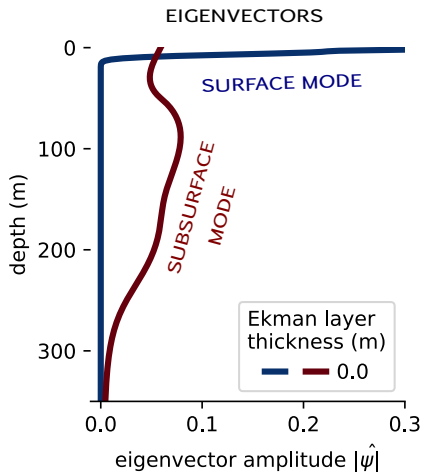
# THE ORIGIN OF MESOSCALE EDDIES

Independent surface- and subsurface-intensified instabilities are identified



# THE ORIGIN OF MESOSCALE EDDIES

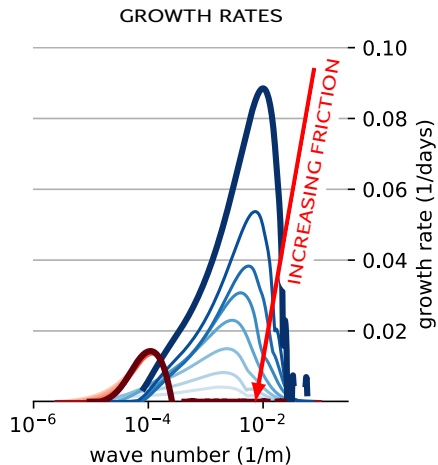
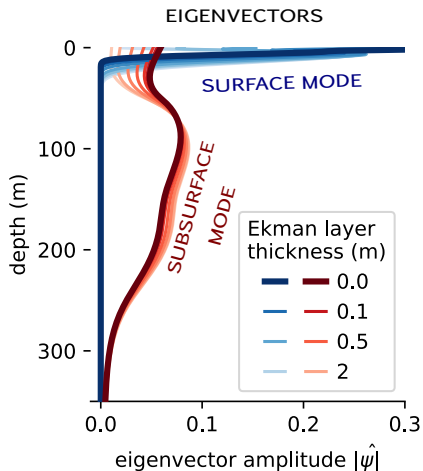
A very close match with observations





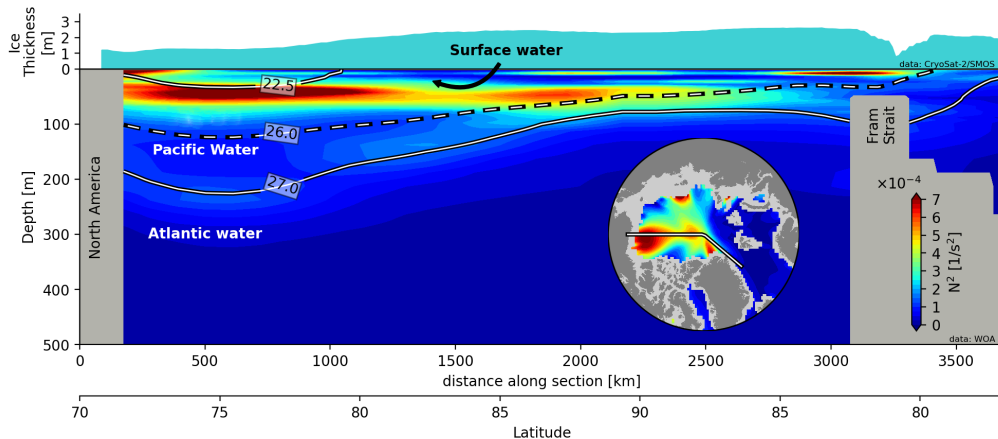
# THE ORIGIN OF MESOSCALE EDDIES

Only surface perturbations are affected by friction



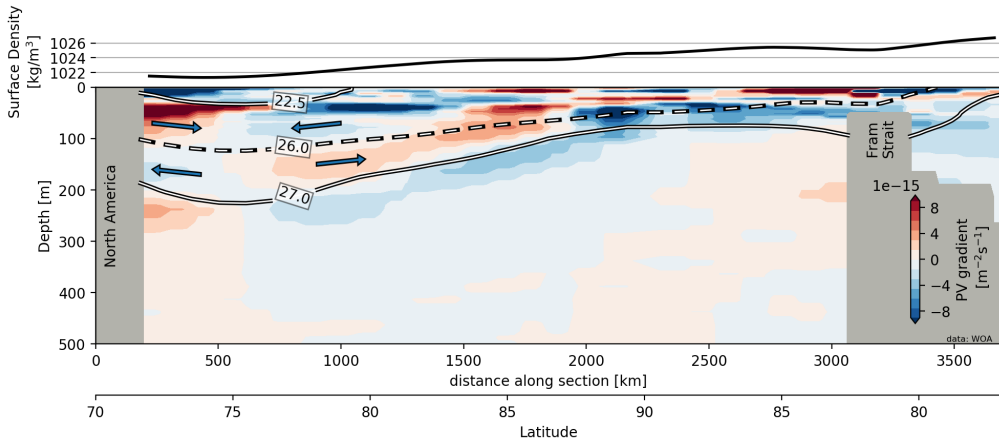
# FROM A MOORING TO THE ENTIRE ARCTIC

Peaks in stratification extend all across the Arctic



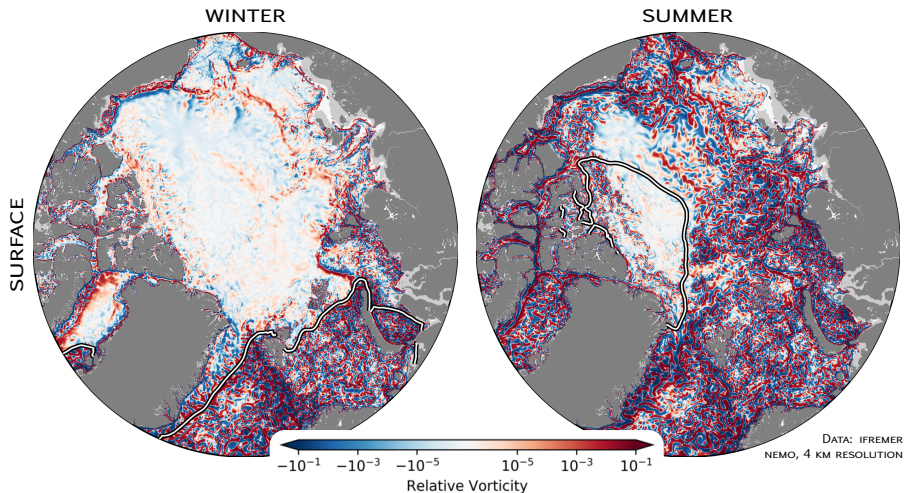
# FROM A MOORING TO THE ENTIRE ARCTIC

Interior PV gradients enable the development of subsurface-intensified turbulence.



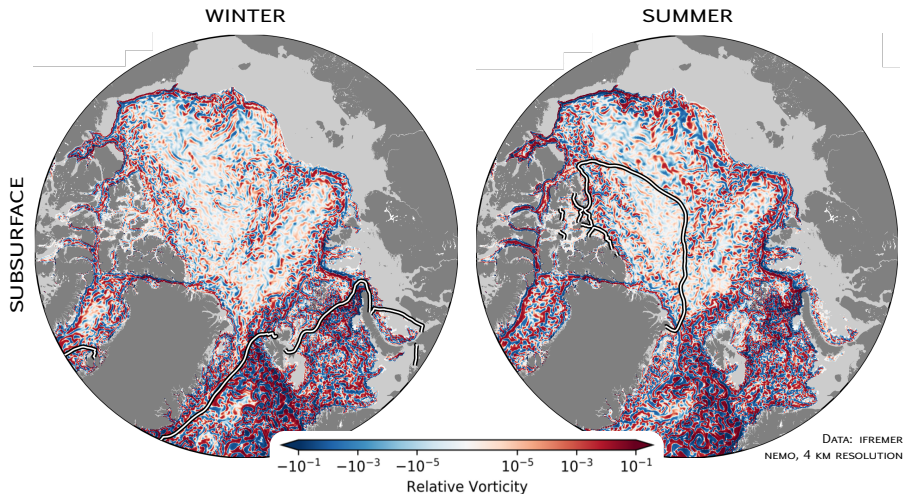
# FROM A MOORING TO THE ENTIRE ARCTIC

High-resolution models support theoretical results



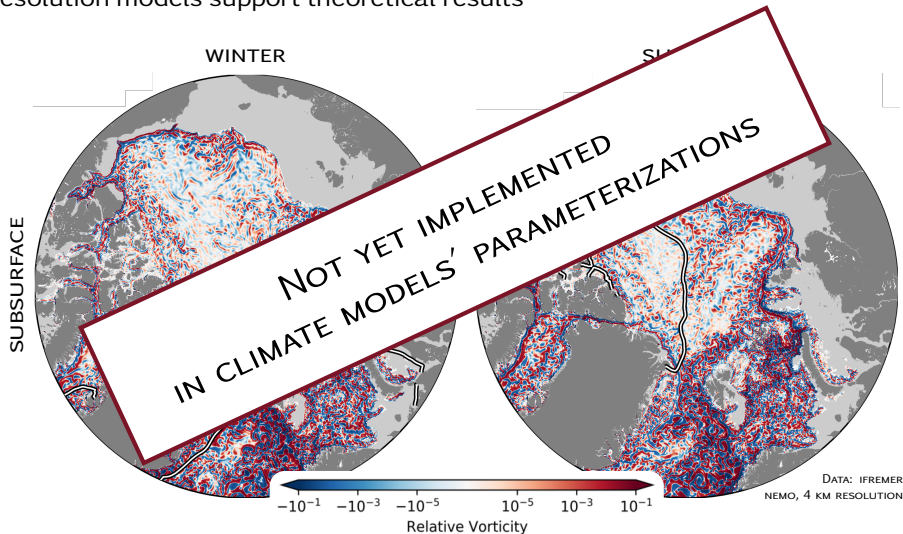
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High-resolution models support theoretical results

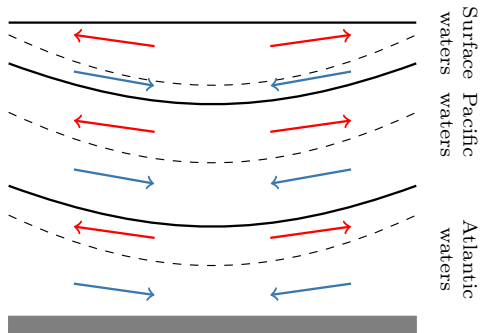


# FROM A MOORING TO THE ENTIRE ARCTIC

High-resolution models support theoretical results



# THE ORIGIN OF MESOSCALE EDDIES — THE SIMPLEST MODEL



## Quasi-geostrophic balance

$$\frac{dq}{dt} = -\mathbf{u} \cdot \nabla Q$$

PV conservation  
(interior)

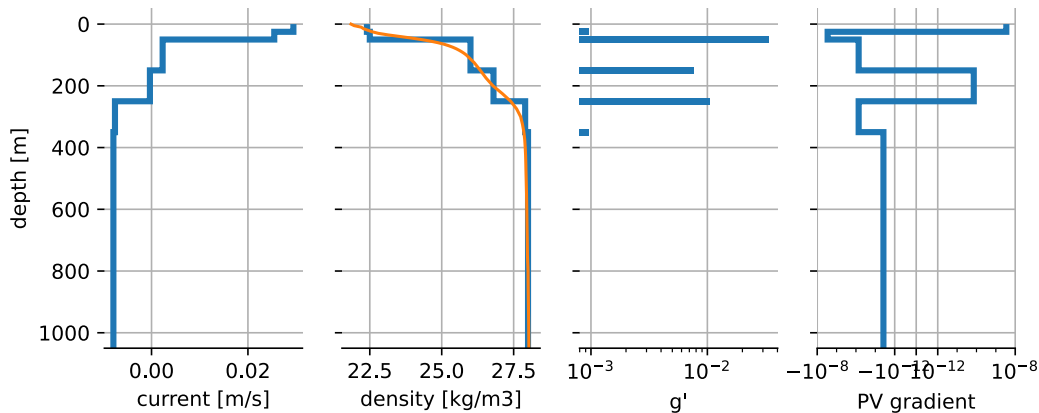
$$\nabla Q = f_0 \frac{\nabla h}{H} = f_0 \frac{d\mathbf{S}}{dz}$$

Background  
PV gradient

Vertical variations of isopycnal slope  
control subsurface eddies.

# UNSTABLE MODES

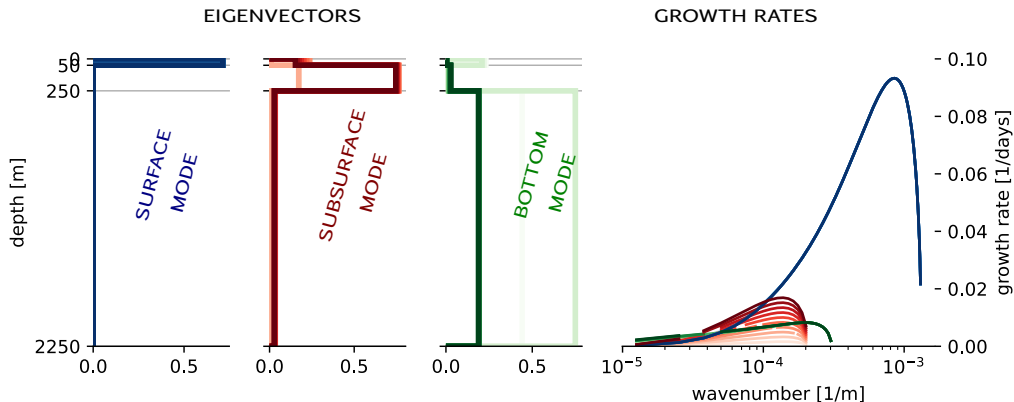
Changes in the Pacific waters' PV gradient only affect the subsurface mode





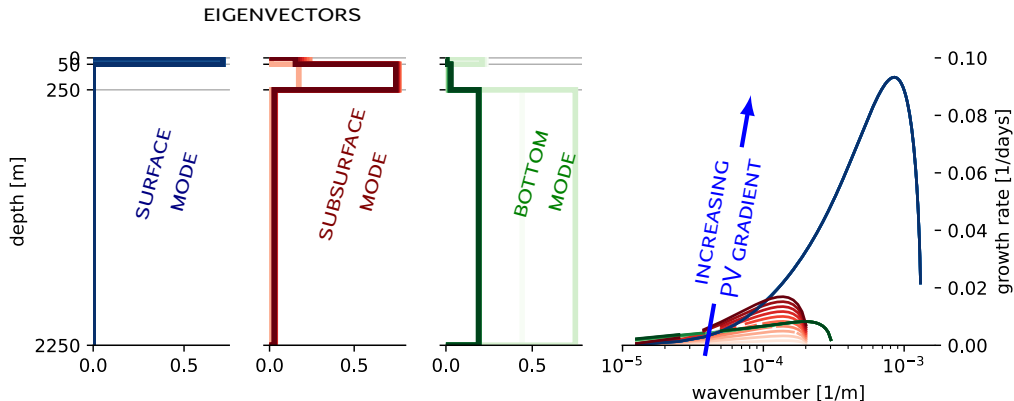
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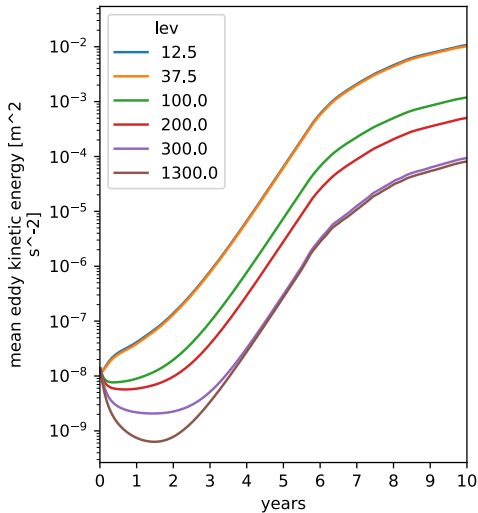
# UNSTABLE MODES

Changes in the Pacific waters' PV gradient only affect the subsurface mode

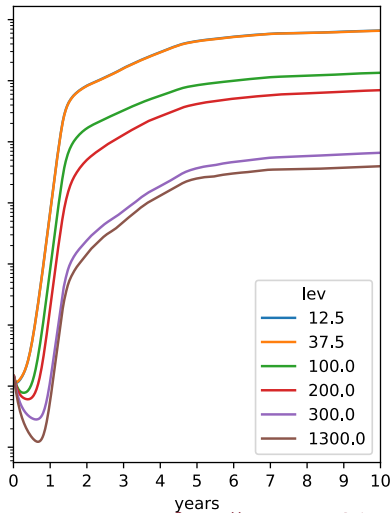


# SIMULATION — EDDY KINETIC ENERGY

## Constant isopycnal slope

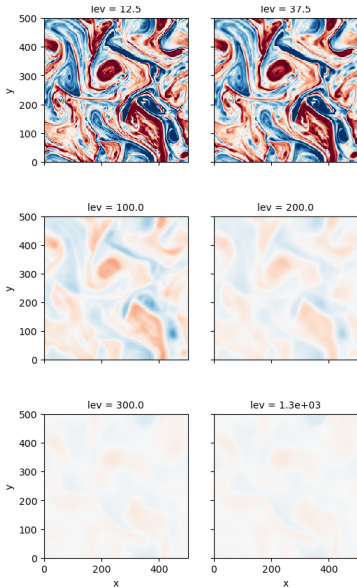


## Varying isopycnal slope

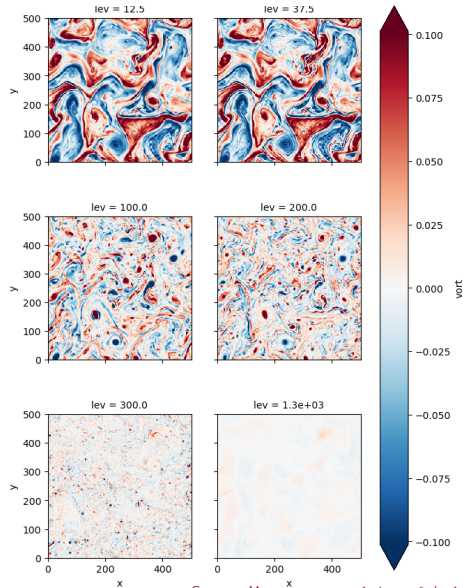


# VORTICITY

## Constant isopycnal slope

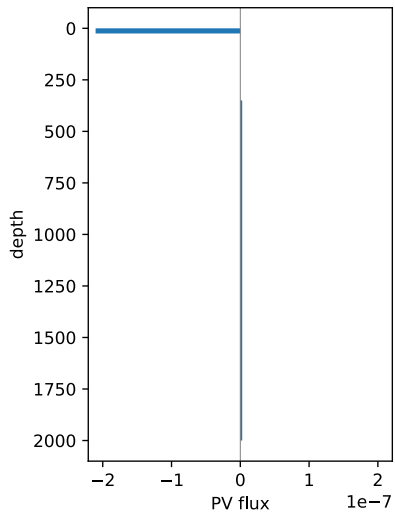


## Varying isopycnal slope

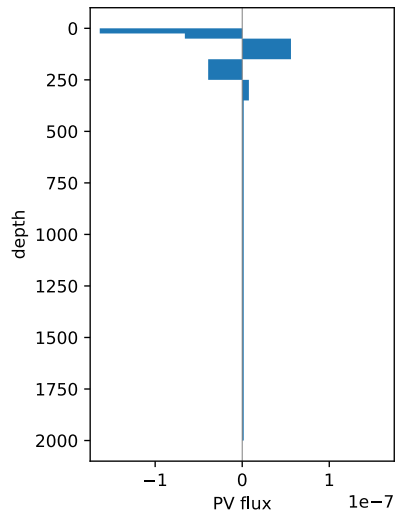


# EDDY FLUXES

## Constant isopycnal slope



## Varying isopycnal slope



# CONCLUSIONS

- ▶ To each water mass, their own eddy system
- ▶ Surface eddies are controlled by the presence of sea ice
- ▶ Subsurface eddies are controlled by internal PV gradients  
... and cannot exist without them!
- ▶ Independent PV fluxes are generated within each water mass

