

# Morphology of Stratocumulus clouds, boundary-layer dynamics and cloud feedback

Florent Brient

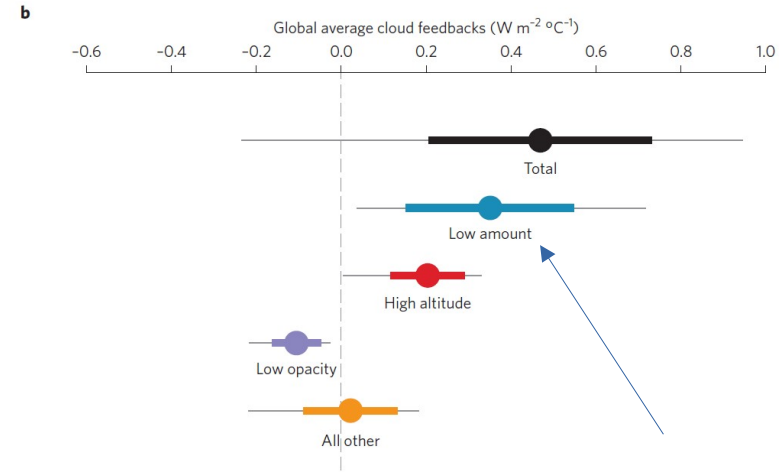
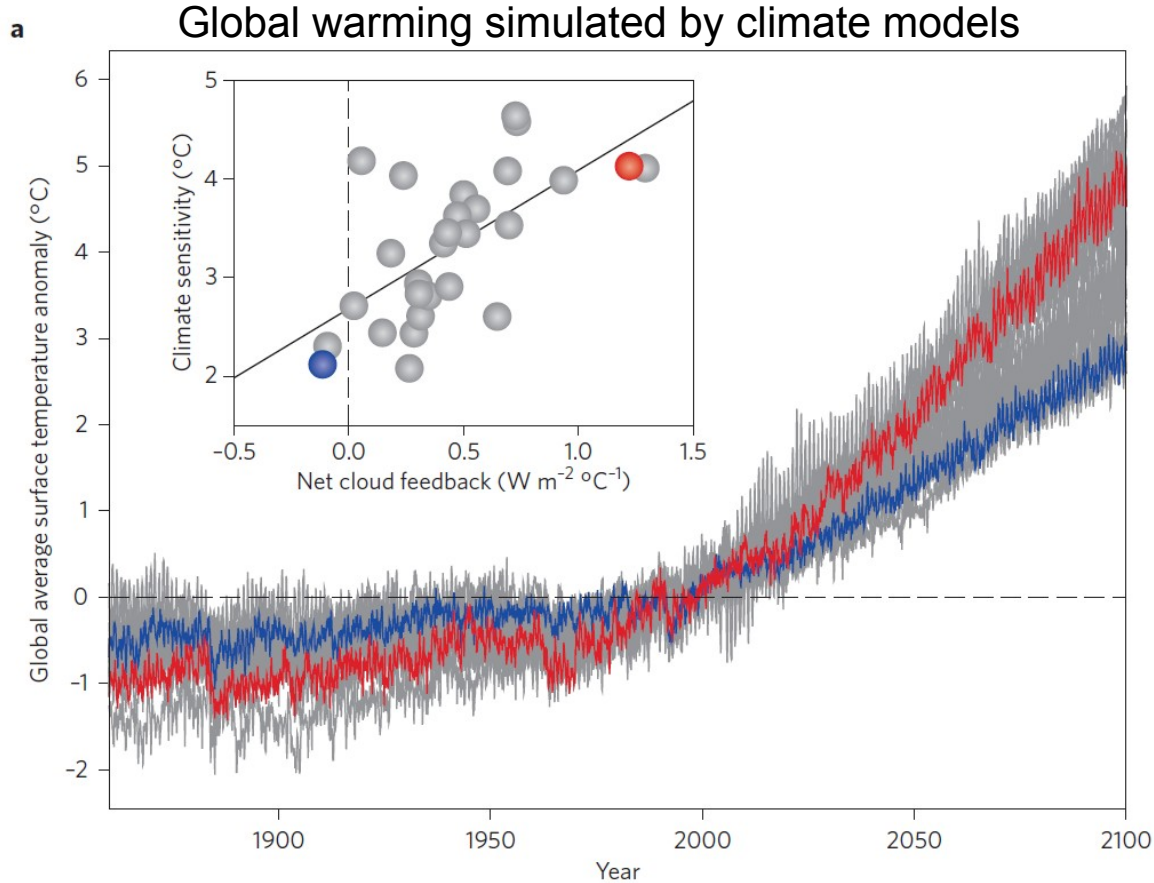
Sorbonne Université - LMD/IPSL

Journées GDR Défis théoriques des sciences du climat

28 mai 2024



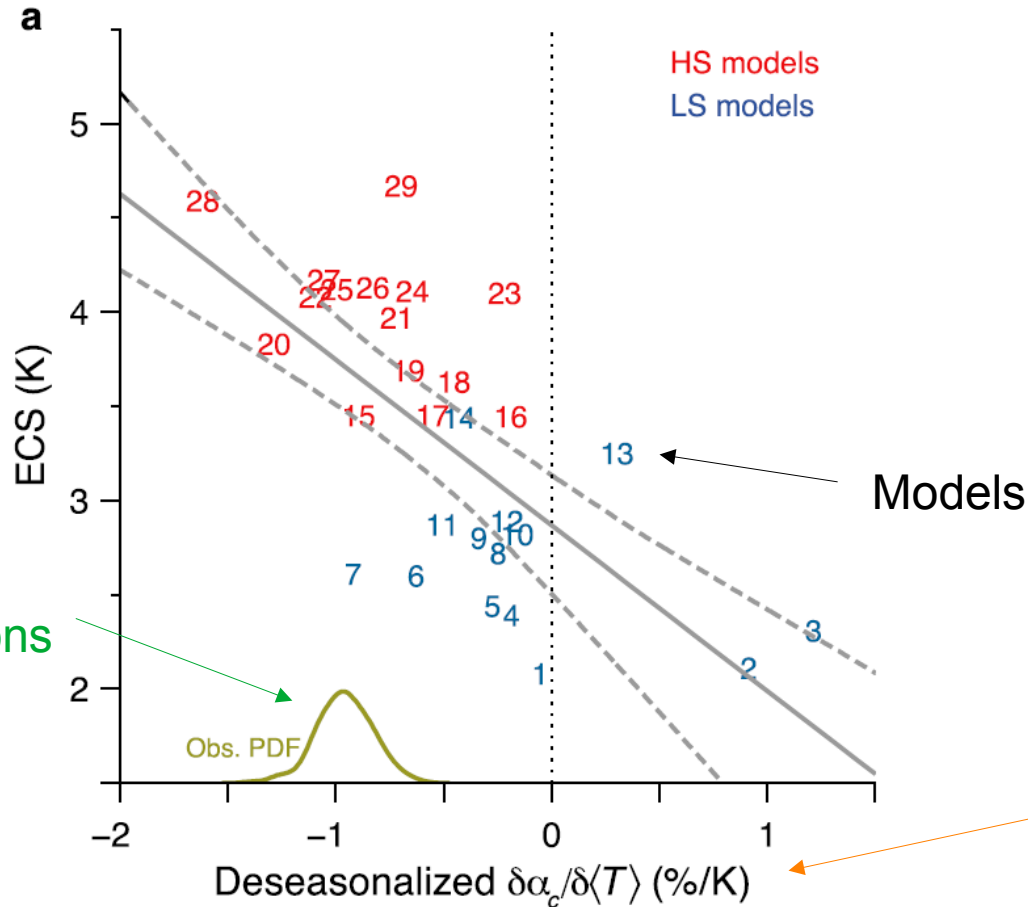
# Clouds of uncertainty



The **low cloud response** remains the most important uncertainty in climate-change projections for a given increase in carbon dioxide concentrations

# Constraining cloud changes

Climate sensitivity



Satellite observations

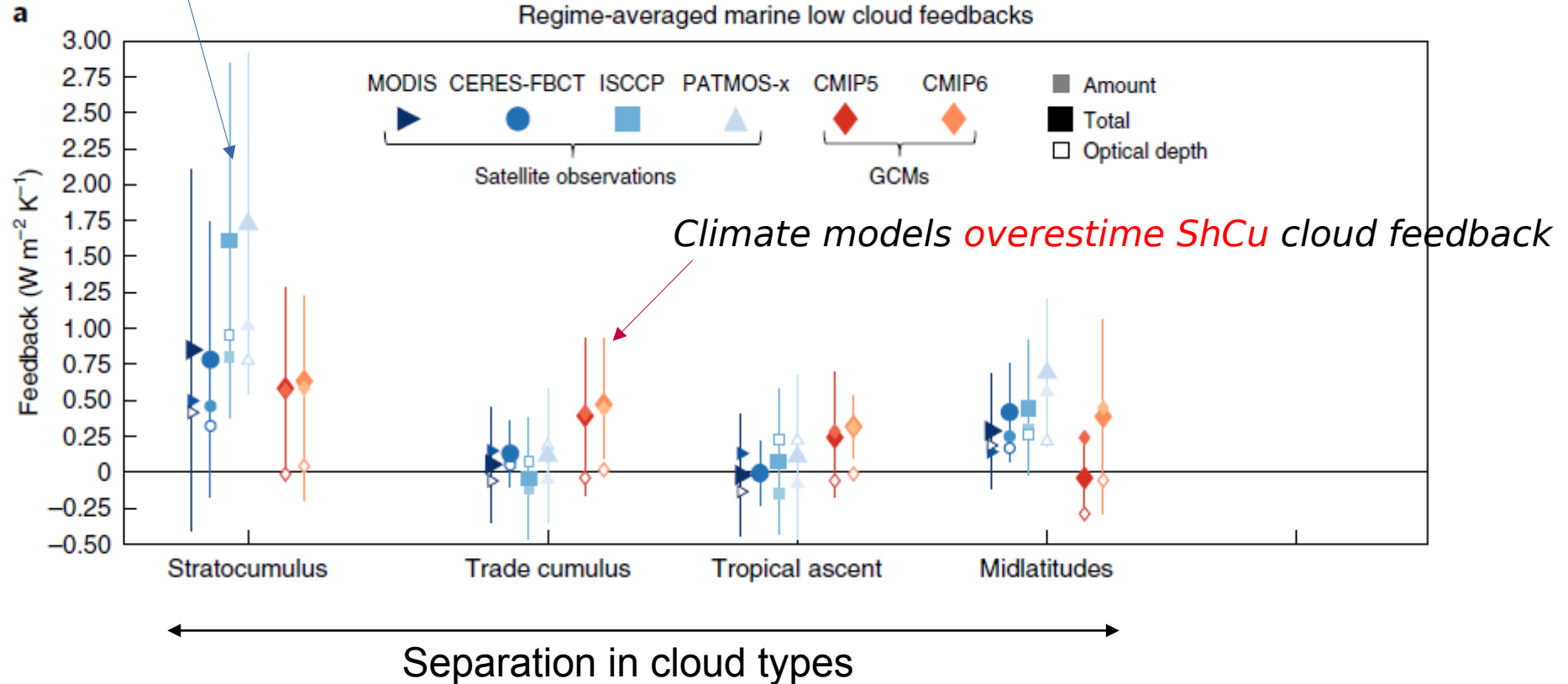
Emergent constraint:  
Climate models that show a present-day **decrease in cloud albedo** with warming and a **high climate sensitivity** seem more realistic

Observable low-cloud variability

*Brient and Schneider (2016)*  
*Brient (2020)*

# Constraining cloud changes

Climate models *underestimate* StCu cloud feedback



Myers et al (2021)



# Constraining cloud changes

## Shallowness of tropical low clouds as a predictor of climate models' response to warming

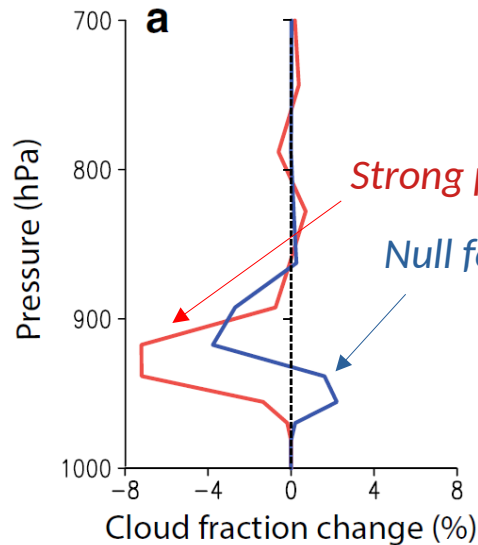
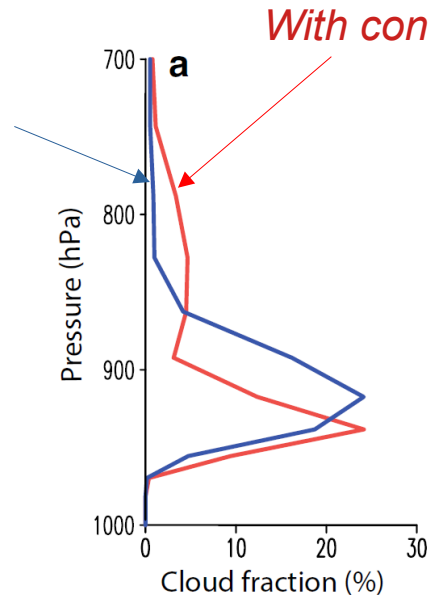
Florent Brient<sup>1</sup> · Tapio Schneider<sup>1,2</sup> · Zhihong Tan<sup>1,2</sup> · Sandrine Bony<sup>3</sup> · Xin Qu<sup>4</sup> · Alex Hall<sup>4</sup>

The shallow cumulus **cloud response** is due to the interaction between **parameterized turbulence and convection**.

The **shallowness** of cloudiness is a signature of **parameterizations**

*Without convection*

IPSL GCM model



Brient et al (2016)

# Constraining cloud changes

Article

## Strong cloud–circulation coupling explains weak trade cumulus feedback

<https://doi.org/10.1038/s41586-022-05364-y> Raphaëla Vogel<sup>1,2</sup>, Anna Lea Albright<sup>1</sup>, Jessica Vial<sup>1</sup>, Geet George<sup>2</sup>, Bjorn Stevens<sup>2</sup> & Sandrine Bony<sup>1</sup>

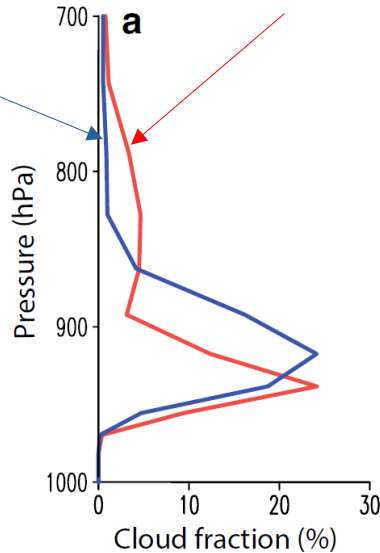
Received: 10 May 2022

The EUREC4A field campaign provides information (data) to constrain mass flux and cloud vertical distribution

Reduction of cloudiness by convection and mass flux is **less realistic** and thus a **strong feedback is unlikely to occur**

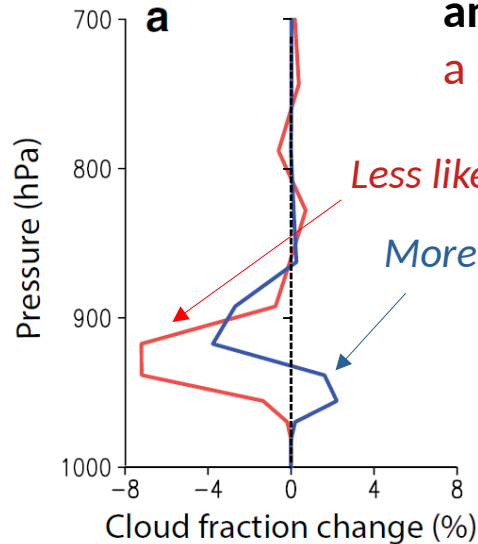
*More realistic*

*Less realistic*



IPSL GCM model

Vogel et al (2022)

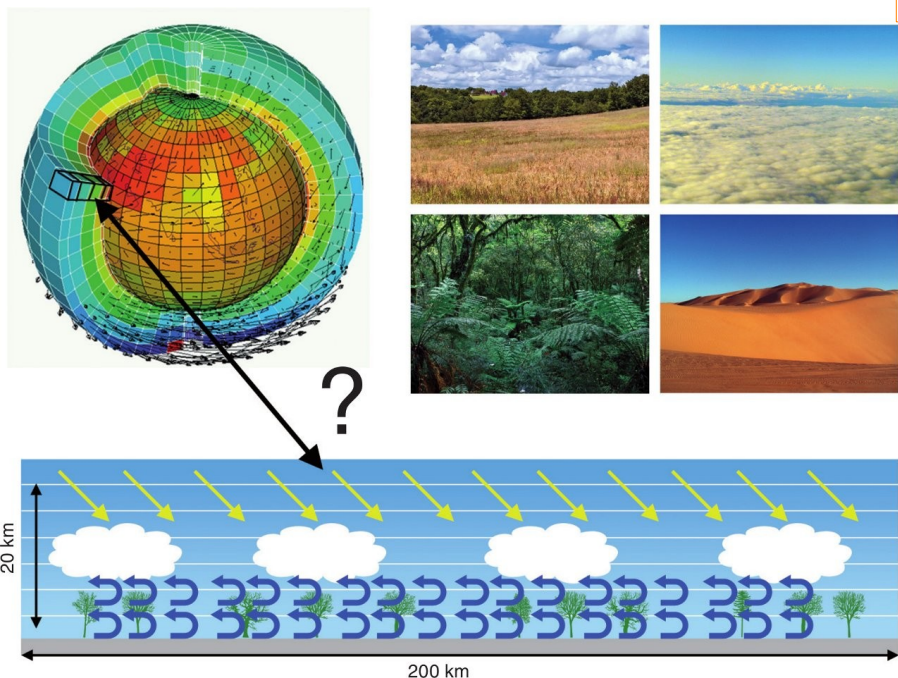


Yet we still **don't** have a clear, robust physical mechanism for **low-cloud feedback**

# Understanding: Modeling and parameterization

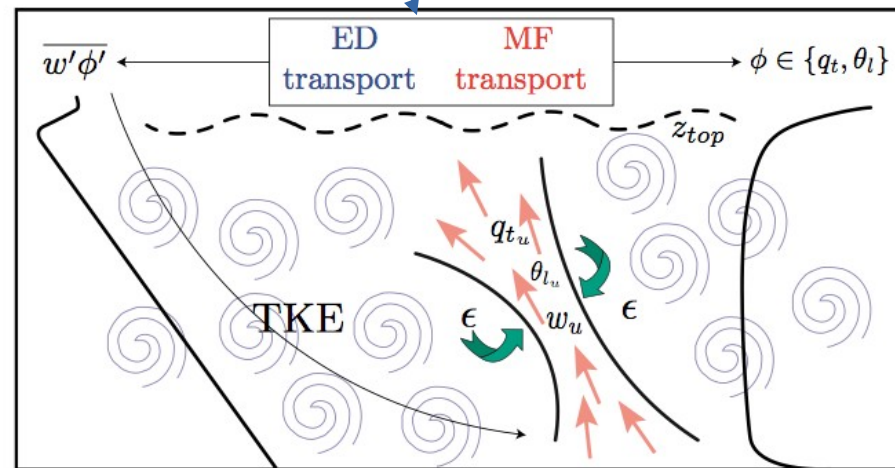
Climate models aim to represent climate variability for different time scales at length scales of around 50-100 km

$$\frac{\partial \bar{\phi}}{\partial t} = -\bar{v} \cdot \nabla \bar{\phi} - \bar{\omega} \frac{\partial \bar{\phi}}{\partial z} - \frac{\partial}{\partial z} \overline{w' \phi'} + \alpha(c - e) + Q_{rad}$$



Resolved

Parameterized



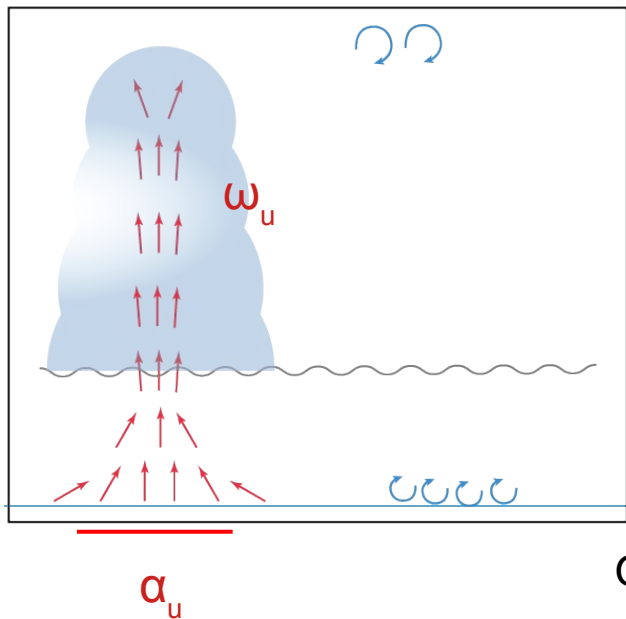
Witte et al (2010)



# Understanding: Modeling and parameterization

Climate models aim to represent climate variability for different time scales at length scales of around 50-100 km

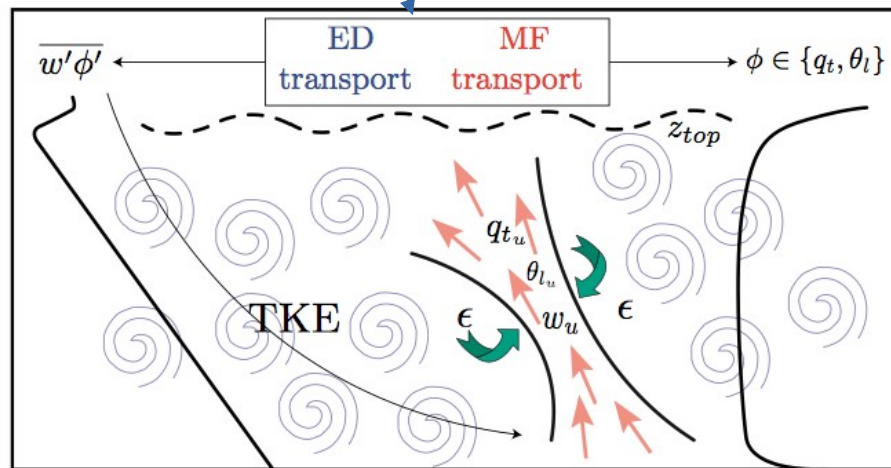
$$\frac{\partial \bar{\phi}}{\partial t} = -\bar{v} \cdot \nabla \bar{\phi} - \bar{\omega} \frac{\partial \bar{\phi}}{\partial z} - \frac{\partial}{\partial z} \overline{\omega' \phi'} + \alpha(c - e) + Q_{rad}$$



Compensating subsidence:  
 $\omega = \alpha_u \omega_u + (1 - \alpha_u) \omega_d$

Resolved

Parameterized



Witte et al (2010)

# High-resolution modeling

Reproducing atmospheric boundary layers to better understanding **coherent structures**, **boundary-layer dynamics** and the **mesoscale organisation**

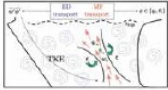




High-Resolution models are the tool for that purpose

The Meso-Nh model is the French mesoscale non-hydrostatic model



Several boundary layers are simulated, three are mostly studied

<http://mesonh.aero.obs-mip.fr/mesonh57>

	Clear-sky	Stratus/Fog	Stratocumulus	St-to-Cu	Cumulus
Warm Cold					
Marine		<i>CGILS (s12)</i>	<b>FIRE</b> <b>DYCOMS</b>	<b>ASTEX</b> <i>CONSTRAIN</i>	<b>RICO</b> <b>BOMEX</b>
Continental	<b>IHOP</b> <b>AYOTTE</b>				<b>ARMCu</b>

## Domain size:

- 12.8x12.8 km<sup>2</sup> (25.6x25.6 km<sup>2</sup> for StCu)
- Double periodic

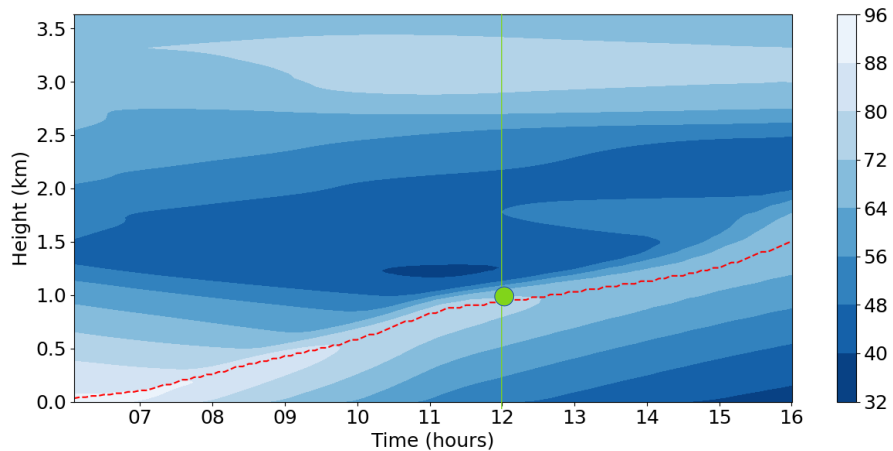
## Resolution:

- $\Delta x = \Delta y = 25\text{m}$  (50m StCu)
- $\Delta z = 25\text{m}$  (10m StCu)
- $\Delta t = 1\text{ sec}$

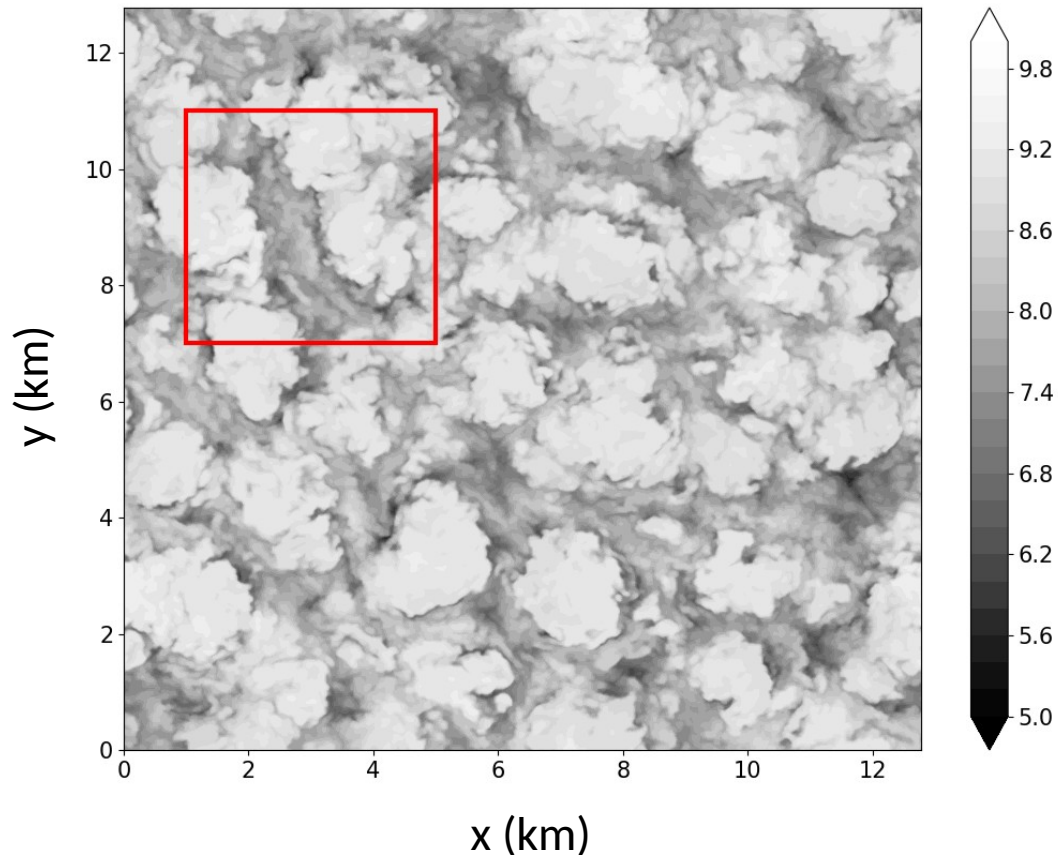
# High-resolution modeling

The clear-sky convective  
boundary layer  
→ no clouds !

Time evolution of averaged Relative Humidity (%)



Cross section of Total humidity (g/kg) at the  
inversion altitude ( $z_i$ )





# Coherent structures

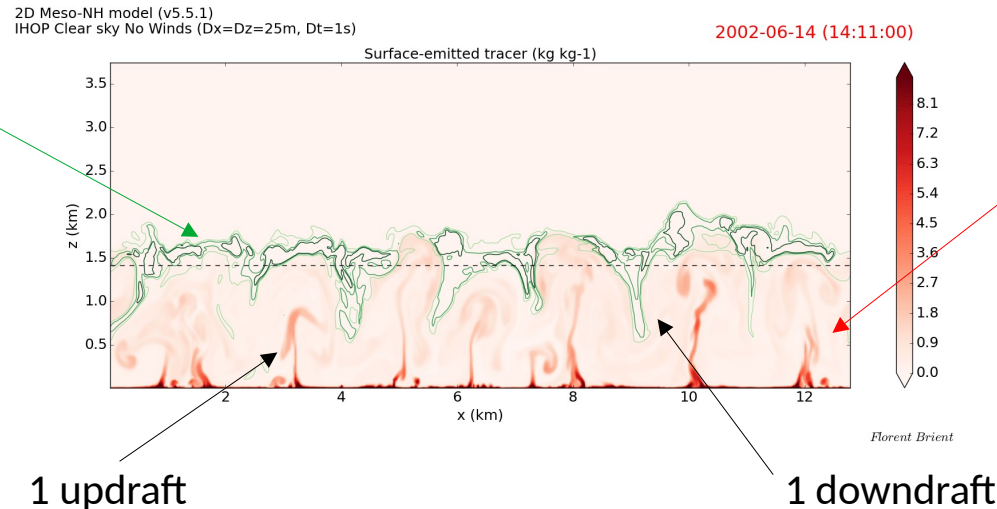
*coherent turbulent structures* = parts of the flow that have logical interconnections and form a unified whole

## Definition:

- **3D Coherent structures** are defined with **passive tracers** emitted at the surface, PBL-top and cloud base
- Ensemble of grid boxes satisfying 2 conditional sampling :  $CS = \{s'(x,y,z) > m * \sigma_s(z)\}$  based on *Couvreur et. al (10)* (with  $s'(x,y,z)$  anomalies of tracer concentrations) and  $CS_w$  for positive/negative vertical velocity
- Object = **3D Contiguous** cells of positive CS (sharing face, edge, corner)
- Selected object = Object with volume **larger than  $V_{min}$**

<https://gitlab.com/tropics/objects>

Cloud tracer emitted at the PBL top



Cloud tracer emitted at the surface

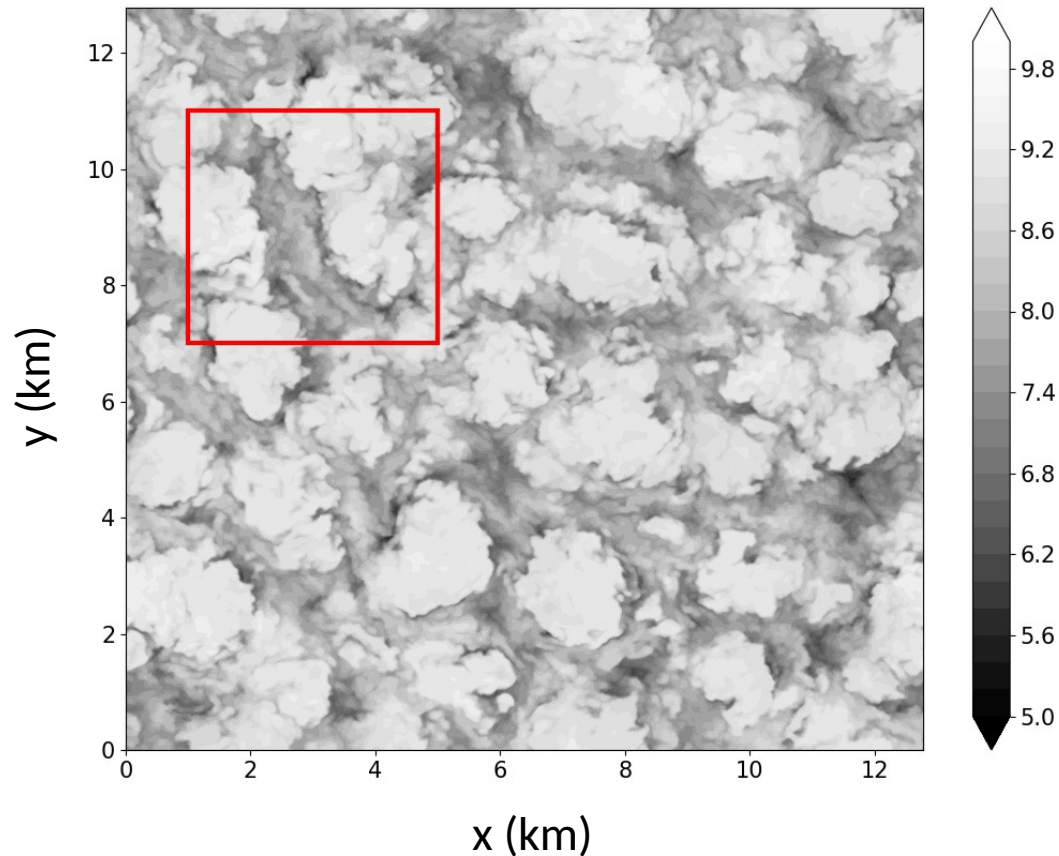
2D simulation of a diurnal cycle of a Dry Convective Boundary Layer (clear-sky)

[\[LINK\]](#)



# High-resolution modeling

The clear-sky convective  
boundary layer  
→ no clouds !



Total humidity (g/kg)  
at the inversion  
altitude ( $z_i$ )

# High-resolution modeling

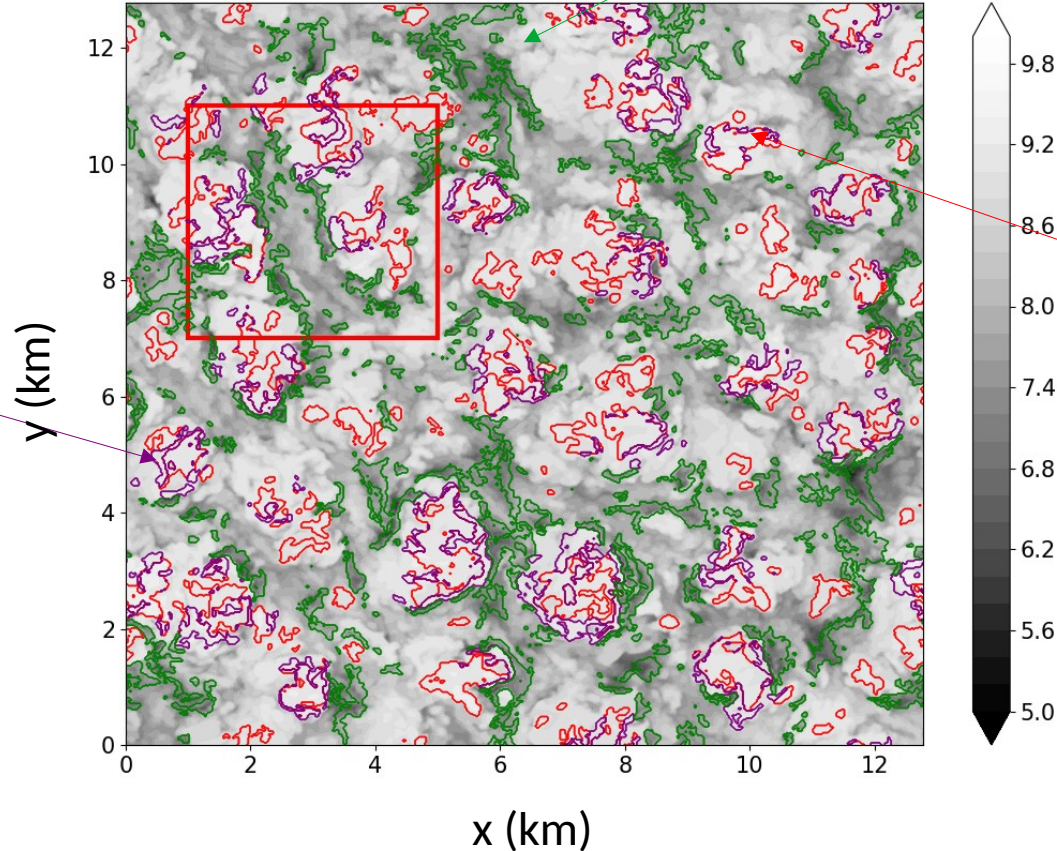
The clear-sky convective boundary layer  
→ no clouds !

Subsiding shells (around updrafts)

Downdraft (between cells)

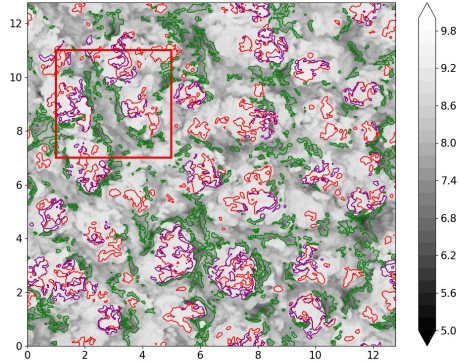
Total humidity (g/kg)  
at the inversion  
altitude ( $z_i$ )

Updraft (in the center  
cells)



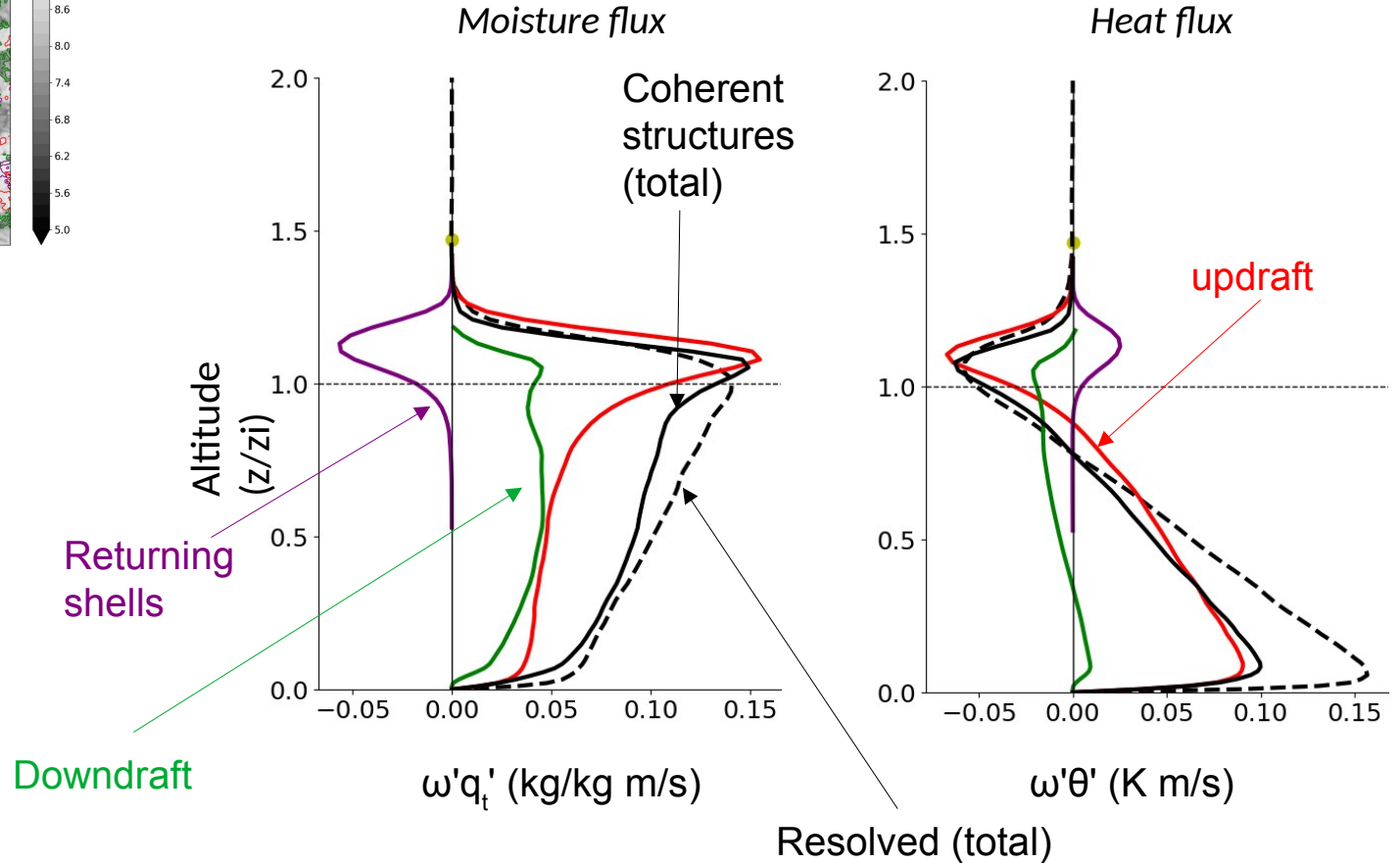


# Coherent structures: Fluxes

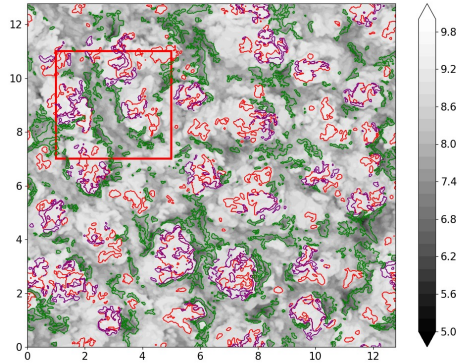


Coherent structures cover **25% of the domain**, but contribute to 70% of resolved **heat** fluxes and 90% of **resolved** moisture fluxes

**Downdrafts** contribute to around **20%** of resolved fluxes



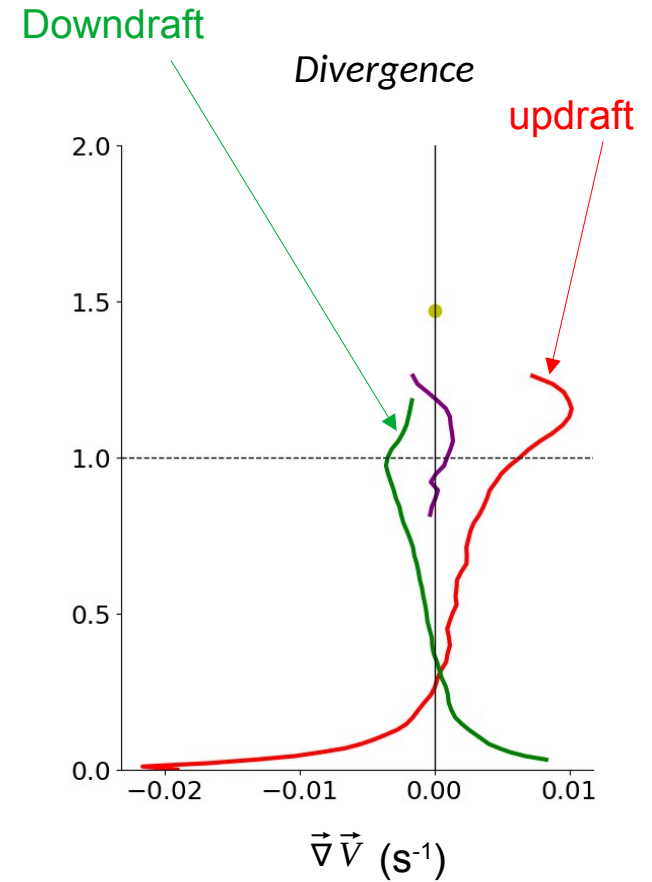
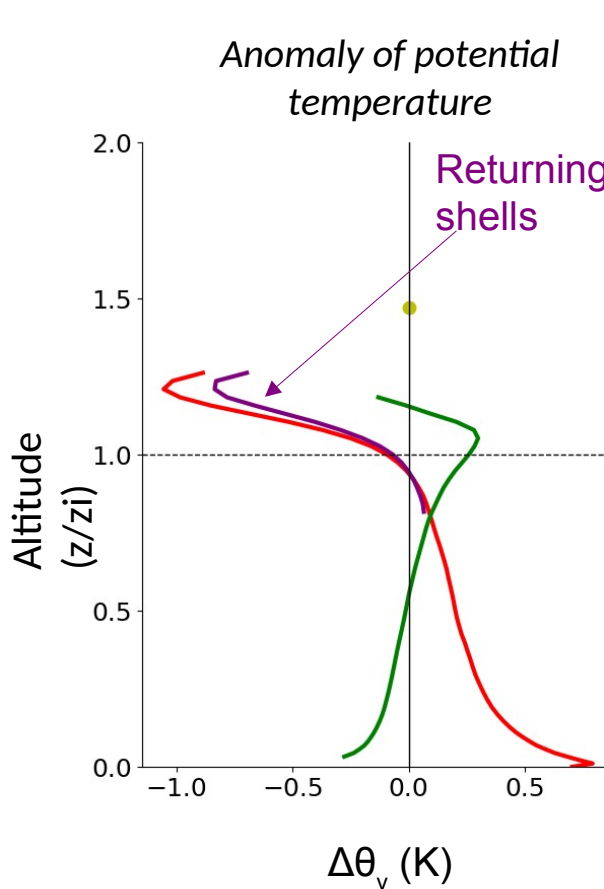
# Coherent structures: Dynamics



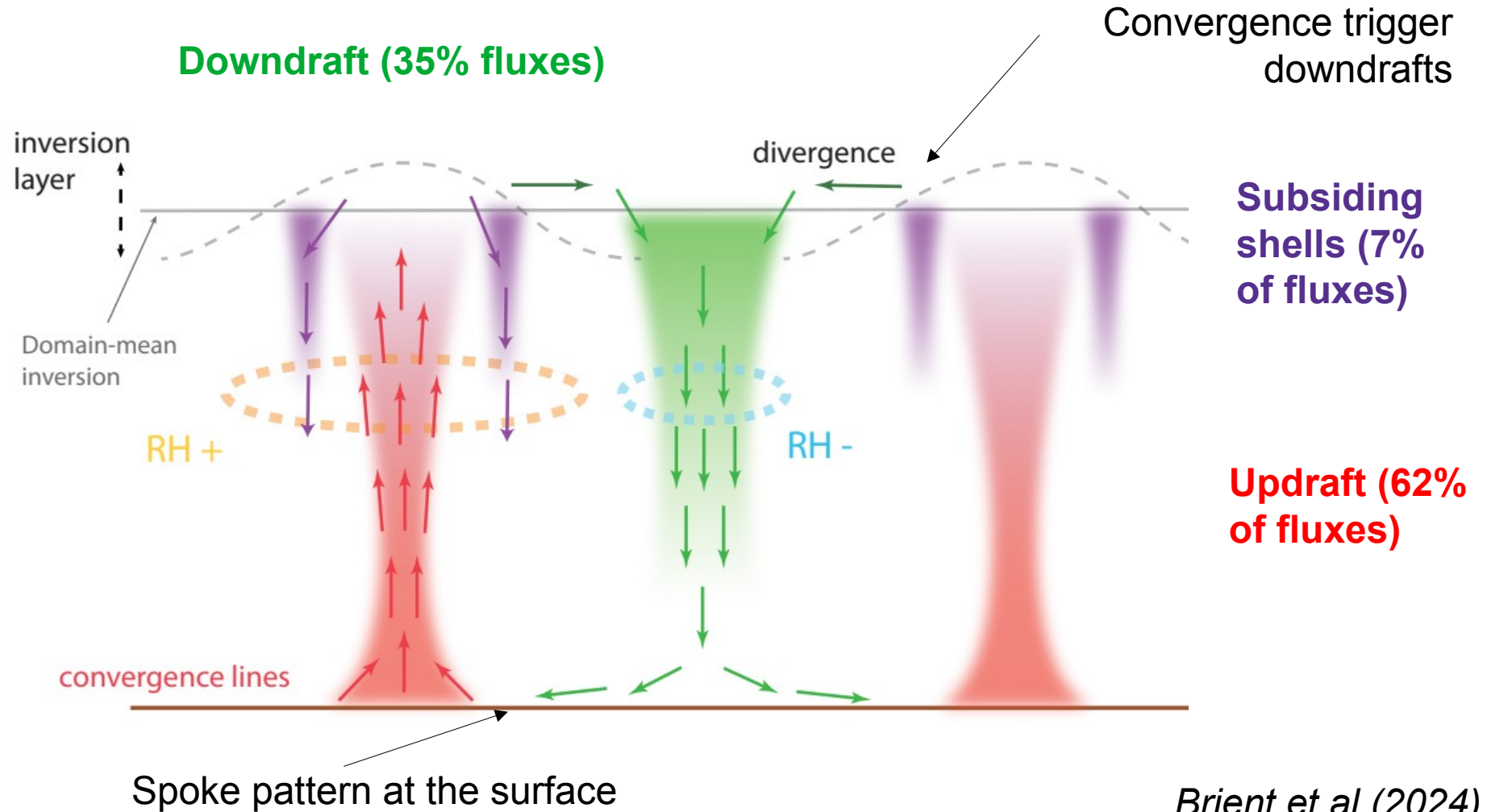
**Updrafts** start positively buoyant at the surface and overshoot at the inversion.

**Returning shells** are located atop the boundary layer, and are similar to updrafts

**Downdrafts** also start **positively buoyancy**, but show **convergence** of air masses  
→ Adiabatic triggering

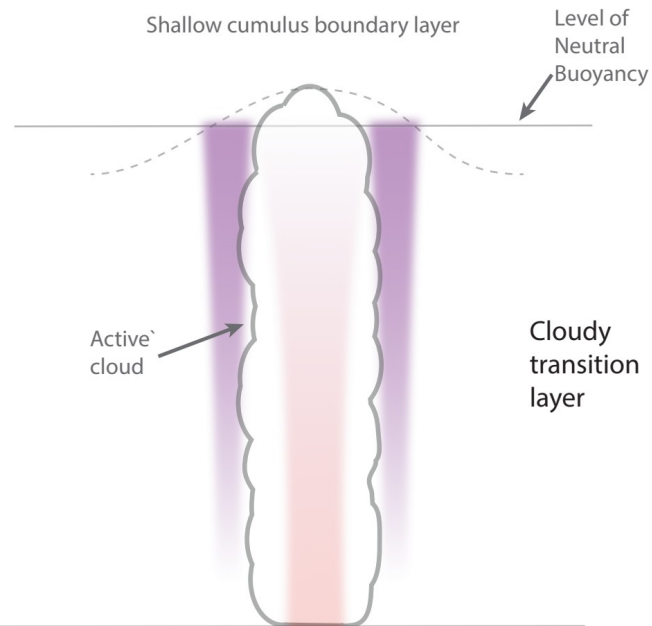
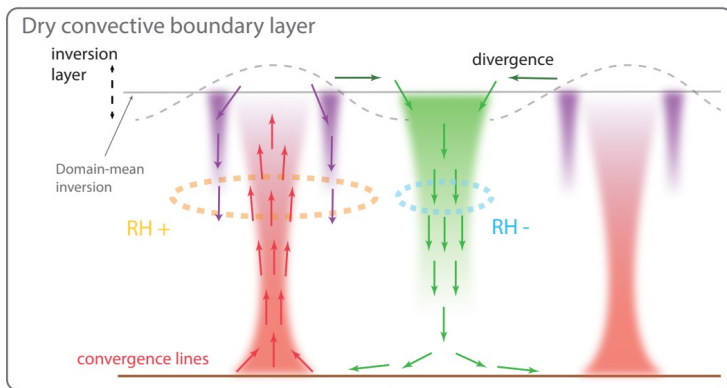


# Schematic of the dry convective boundary layer

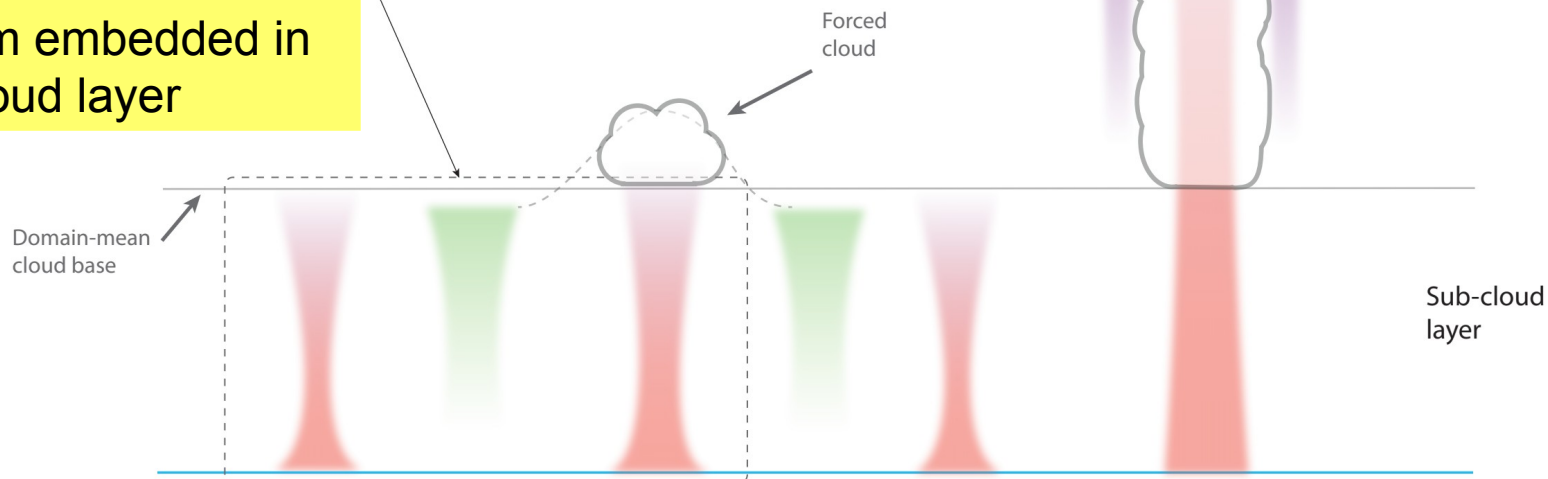




# Schematic of the cumulus boundary layer



Mechanism embedded in the sub-cloud layer



Spoke pattern at the surface

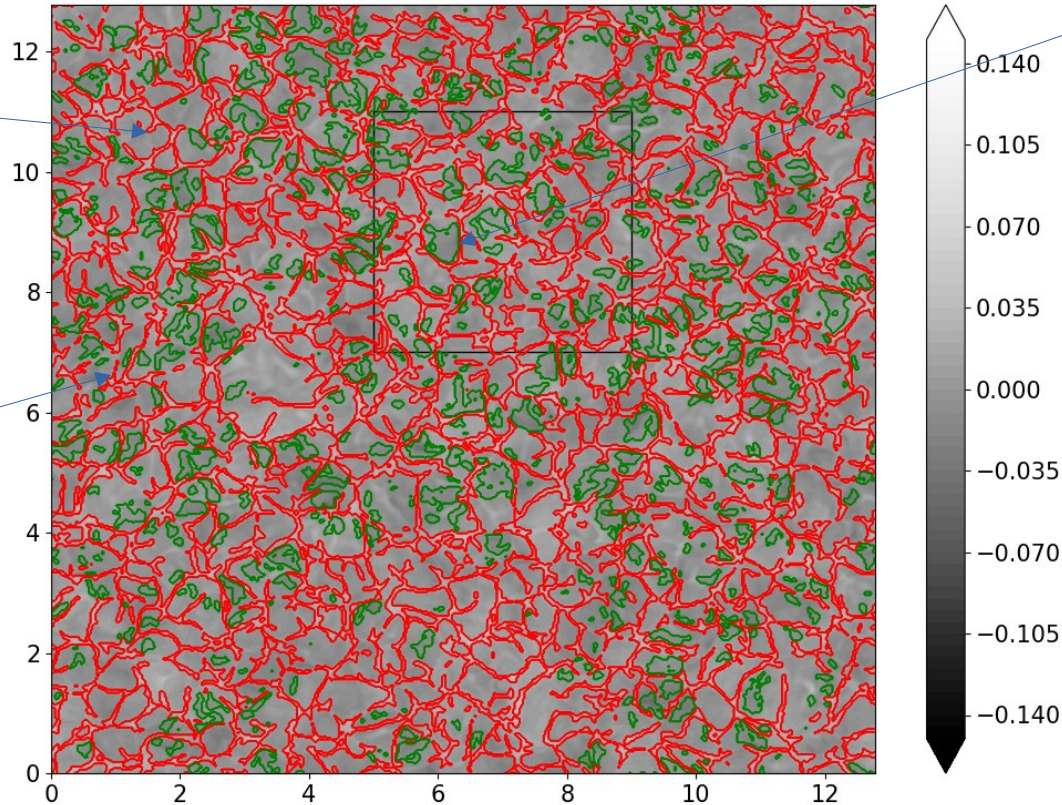
# Spoke pattern at the surface

Cumulus simulation  
(BOMEX)

Anomaly of relative humidity at  $0.1z_i$

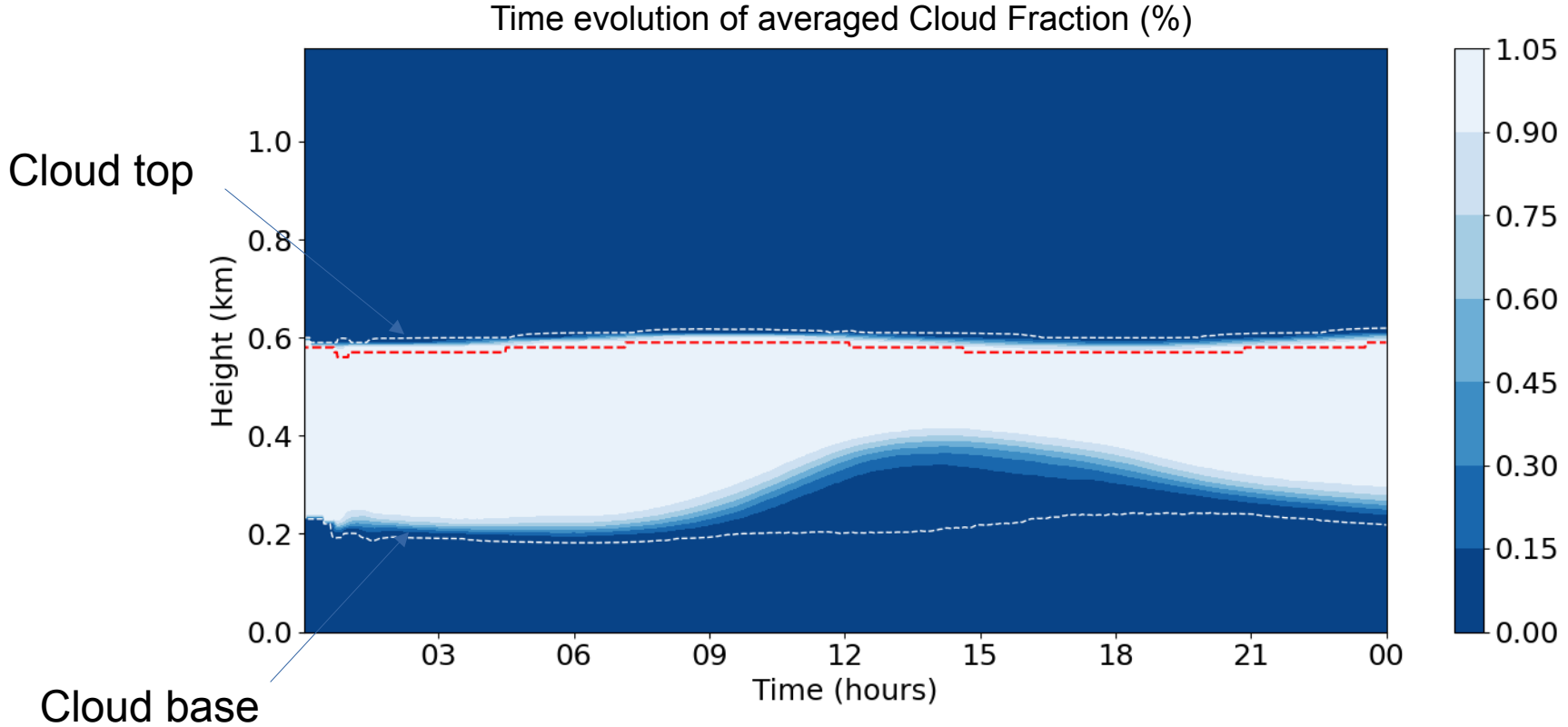
Spoke  
pattern at  
the surface

Convergence  
lines by  
updrafts



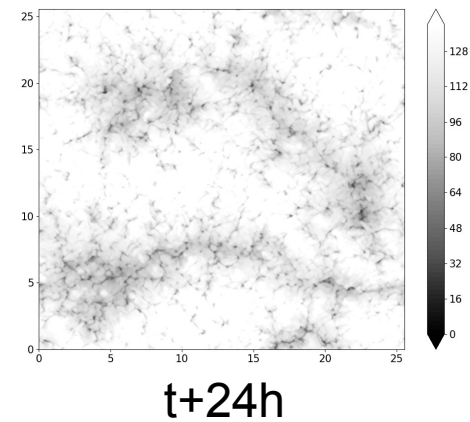
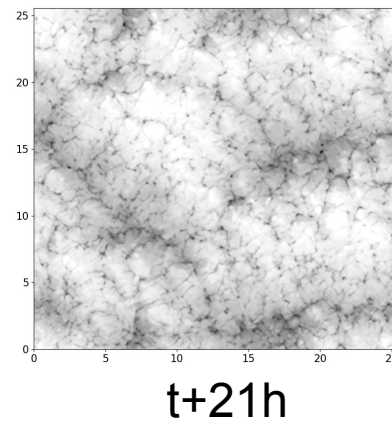
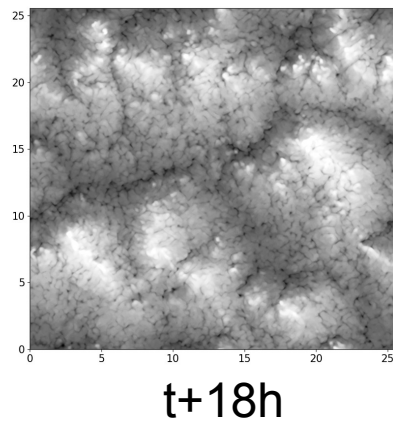
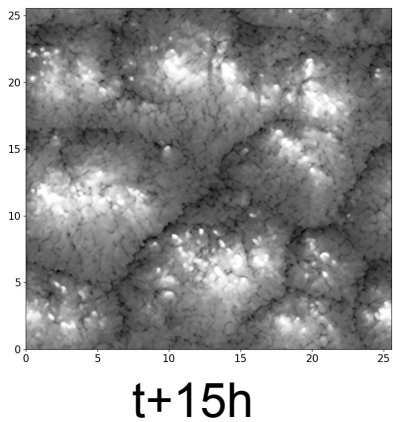
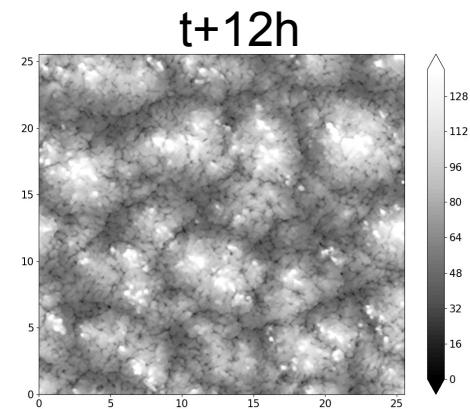
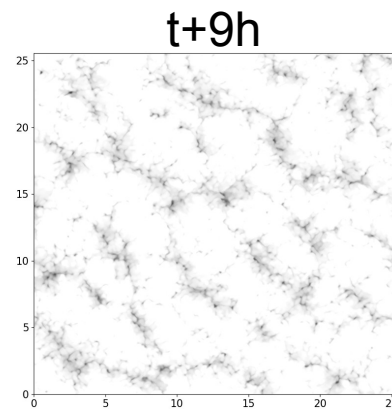
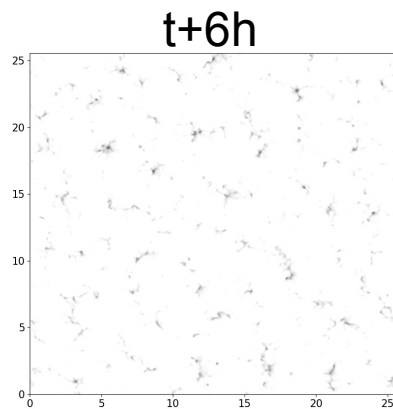
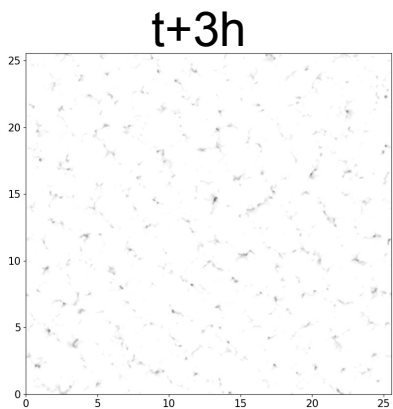
Downdrafts in  
the middle of  
updrafts' circles

# What about stratocumulus?



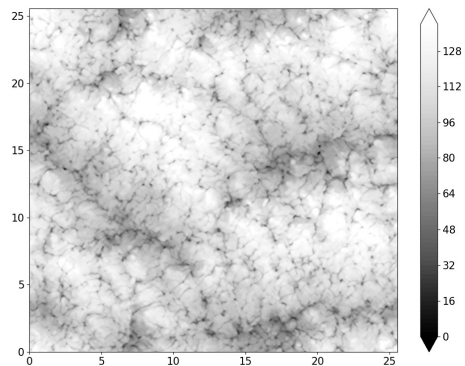
# What about stratocumulus?

Liquid Water Path ( $\text{g}/\text{m}^2$ )





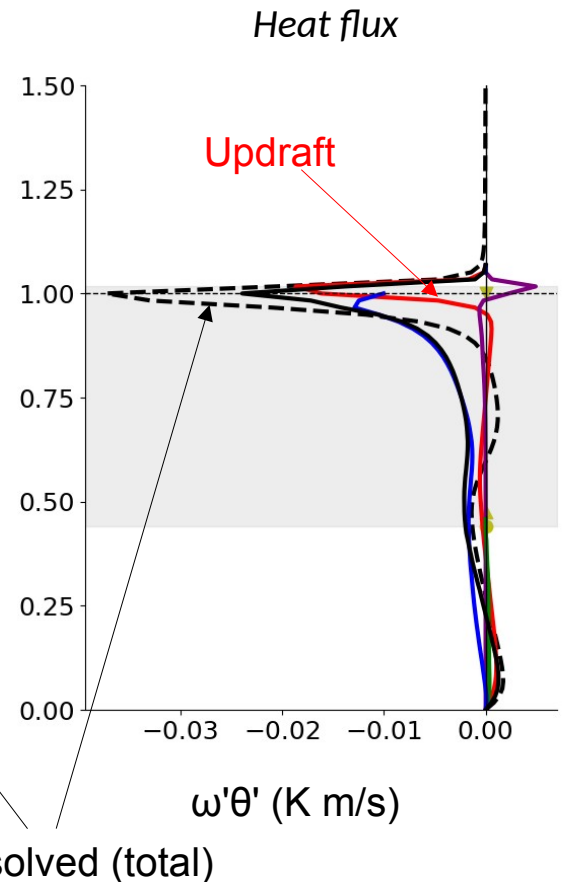
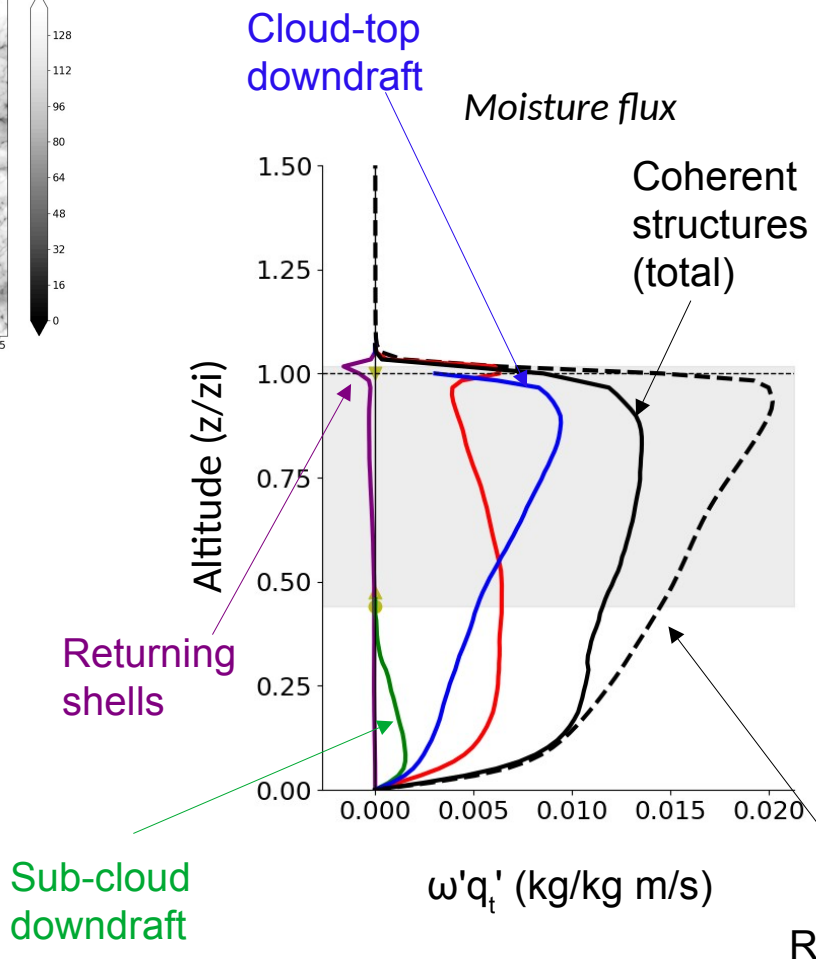
Nighttime ( $t+21h$ )



# Coherent structures: Fluxes

Coherent structures cover **27% of the domain**, but contributes to **78%** of resolved moisture fluxes

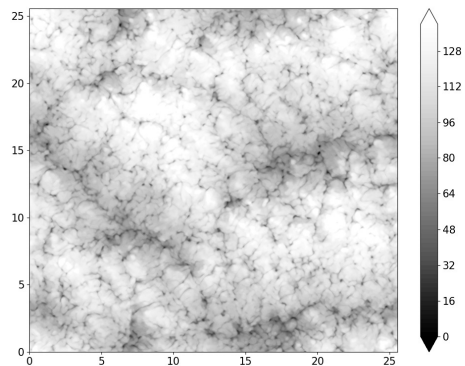
**Cloud-top downdrafts** to around **40%** of resolved fluxes



Resolved (total)

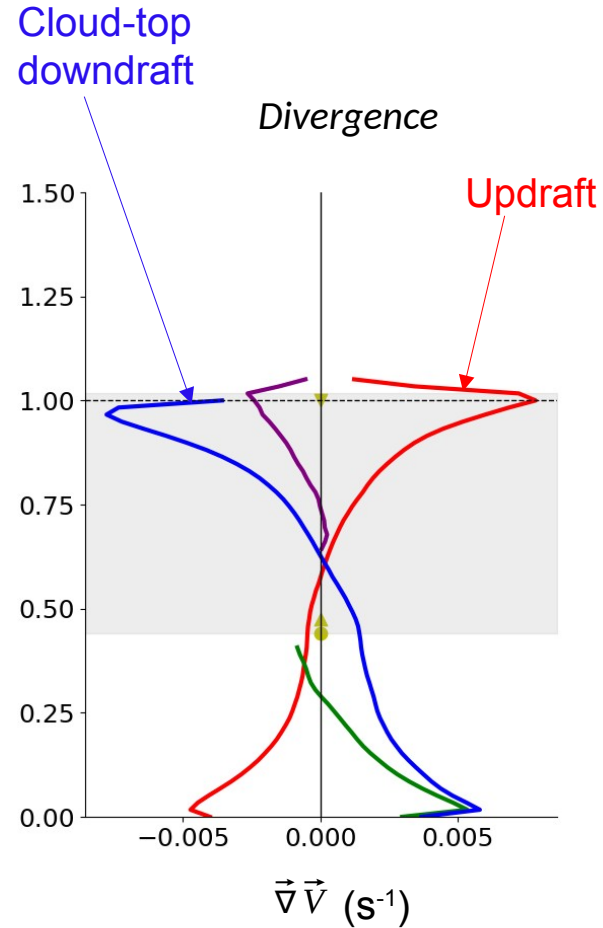
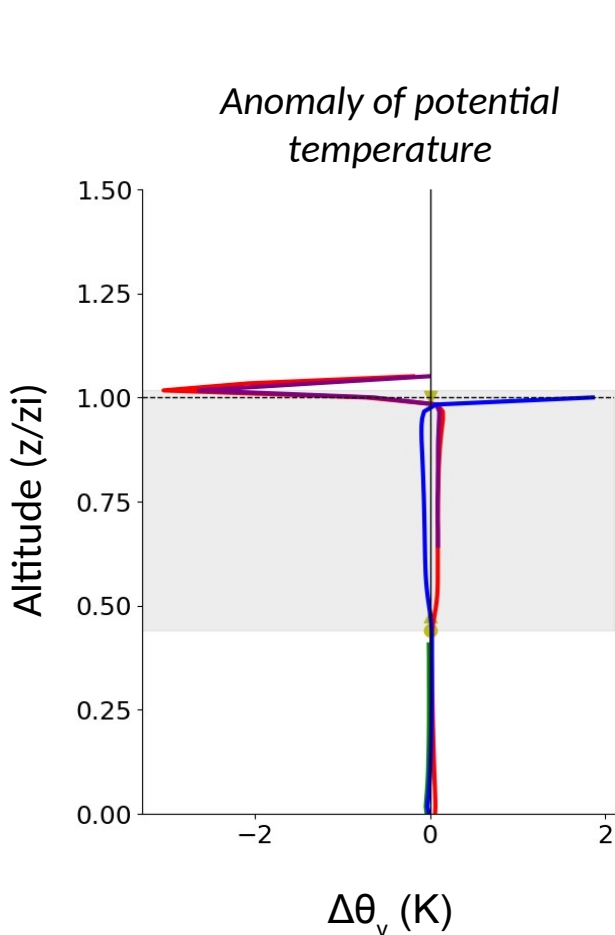
# Coherent structures: Dynamics

Nighttime ( $t+21h$ )

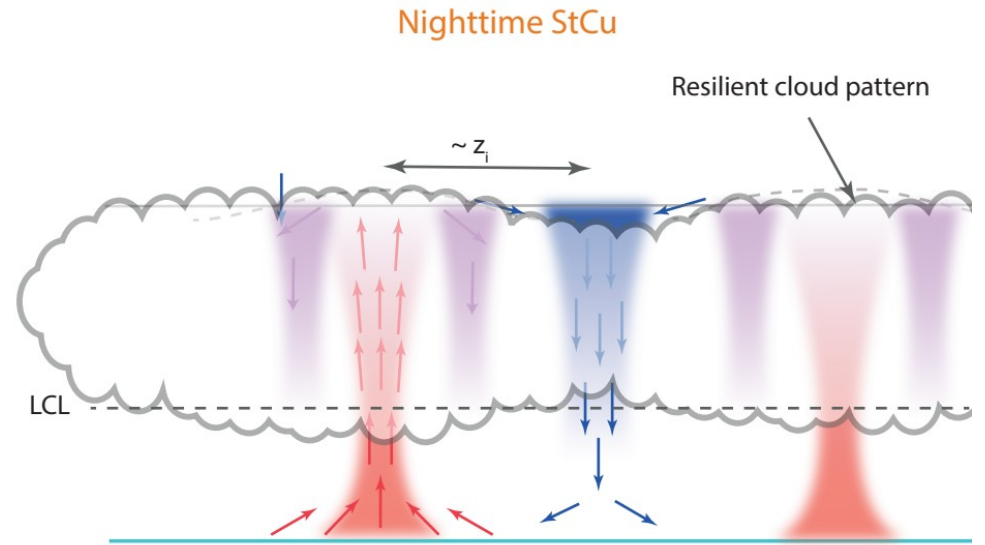


**Updrafts** and **cloud-top downdrafts** have opposite characteristics

Despite strong radiative cooling, **cloud-top downdrafts** start **positively buoyancy** and undergo **convergence** of air masses  
→ Similarities with the dry convective boundary layer !



# Schematic of the stratocumulus boundary layer



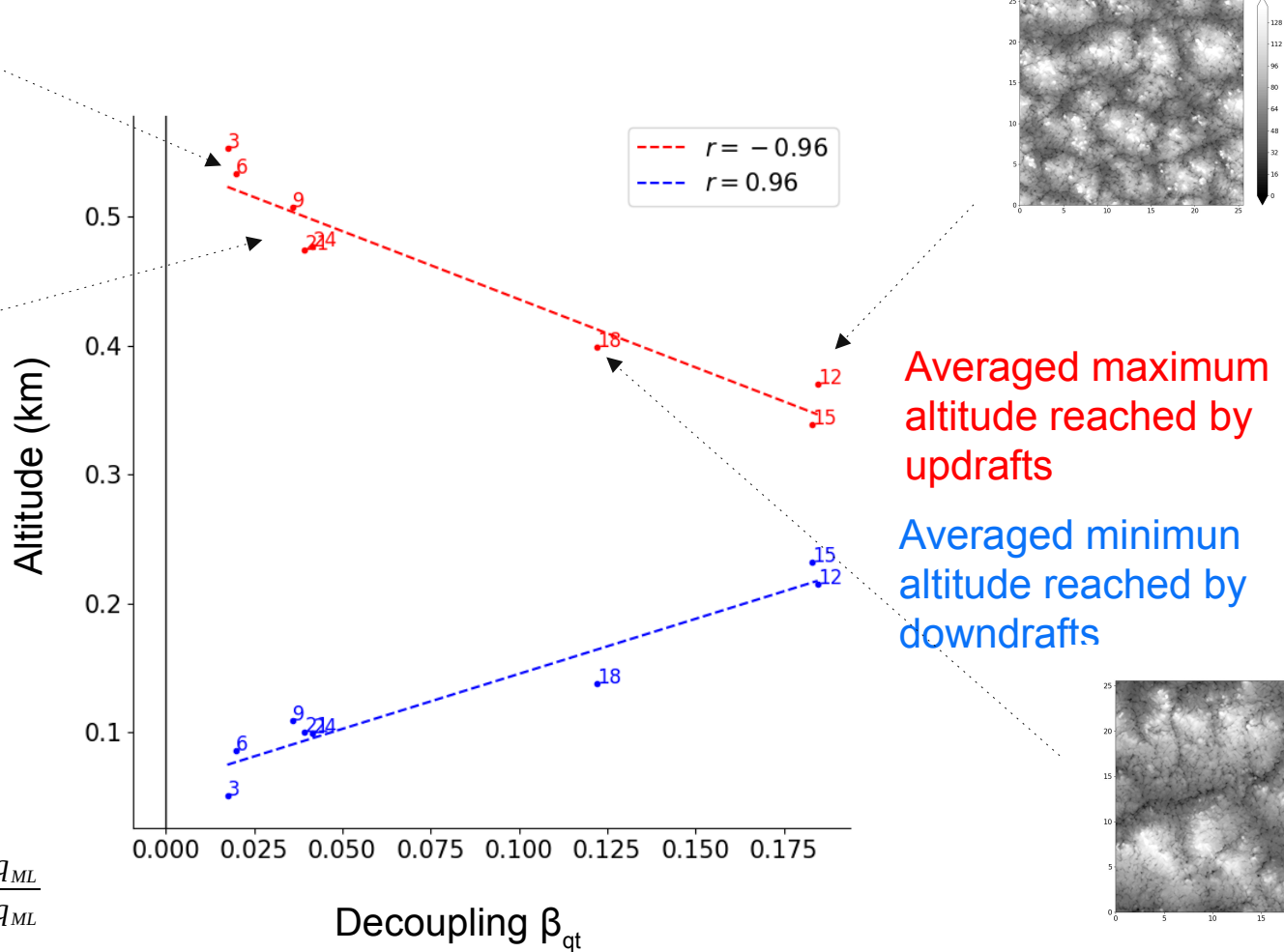
## At nighttime

- Boundary layer is coupled
- **Downdrafts contribute to 80% of fluxes**
-

# Schematic of the stratocumulus boundary layer

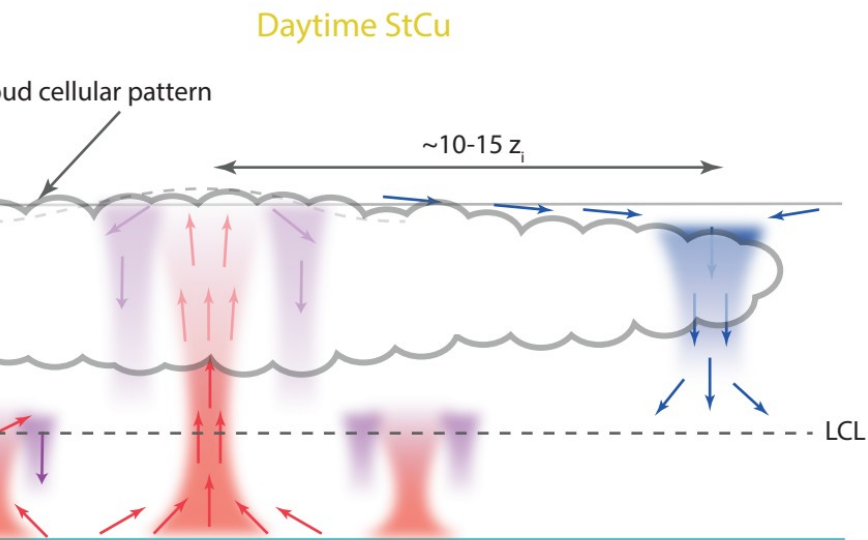
Decoupling index:  
simple quantification  
of stabilization  
(0 = mixed layer)

$$\frac{q_{cld} - q_{ML}}{q_{inv} - q_{ML}}$$



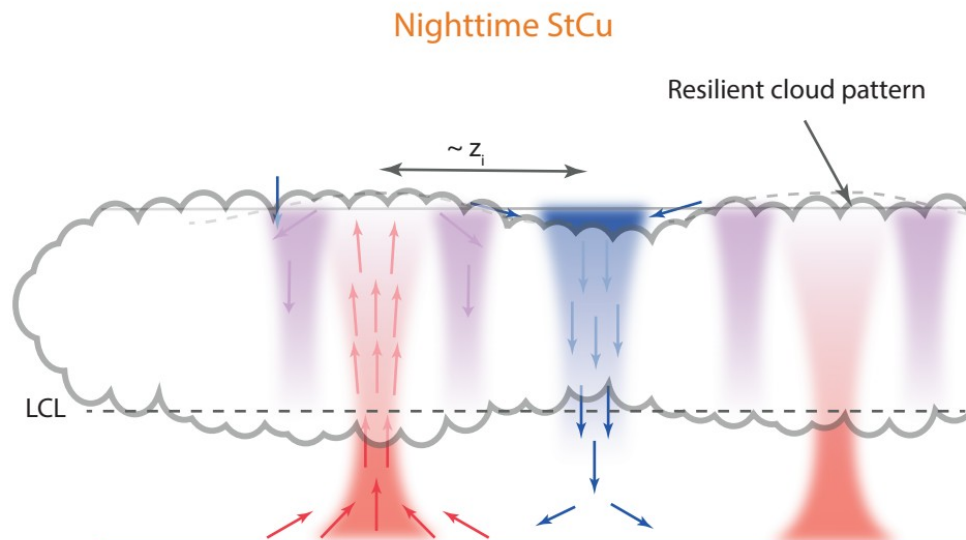


# Schematic of the stratocumulus boundary layer



## At daytime

- Most updrafts are located at the center of the cells, most downdrafts at their surroundings
- Updraft contribute to 50% of fluxes.
- **Decoupling** reduces links between surface and cloud top
- **Aspect ratio of 10-30**



## At nighttime

- Boundary layer is coupled
- **Downdrafts contribute to 80% of fluxes**
- Resilient cloud pattern of the daytime organization

# Intermediate conclusions and remaining questions

## Conclusions

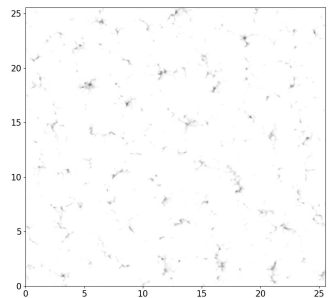
- Passive tracer analysis is really efficient to identify and study **coherent structures**, which contribute to **80% of resolved fluxes** while covering only 25 % of the domain
- Downdrafts are **adiabatically** triggered in all boundary layers. Negative buoyancy is enhanced by radiative/evaporative cooling in stratocumulus
- **Interaction between updrafts and downdrafts** shape the boundary layer organisation

## Questions to go further

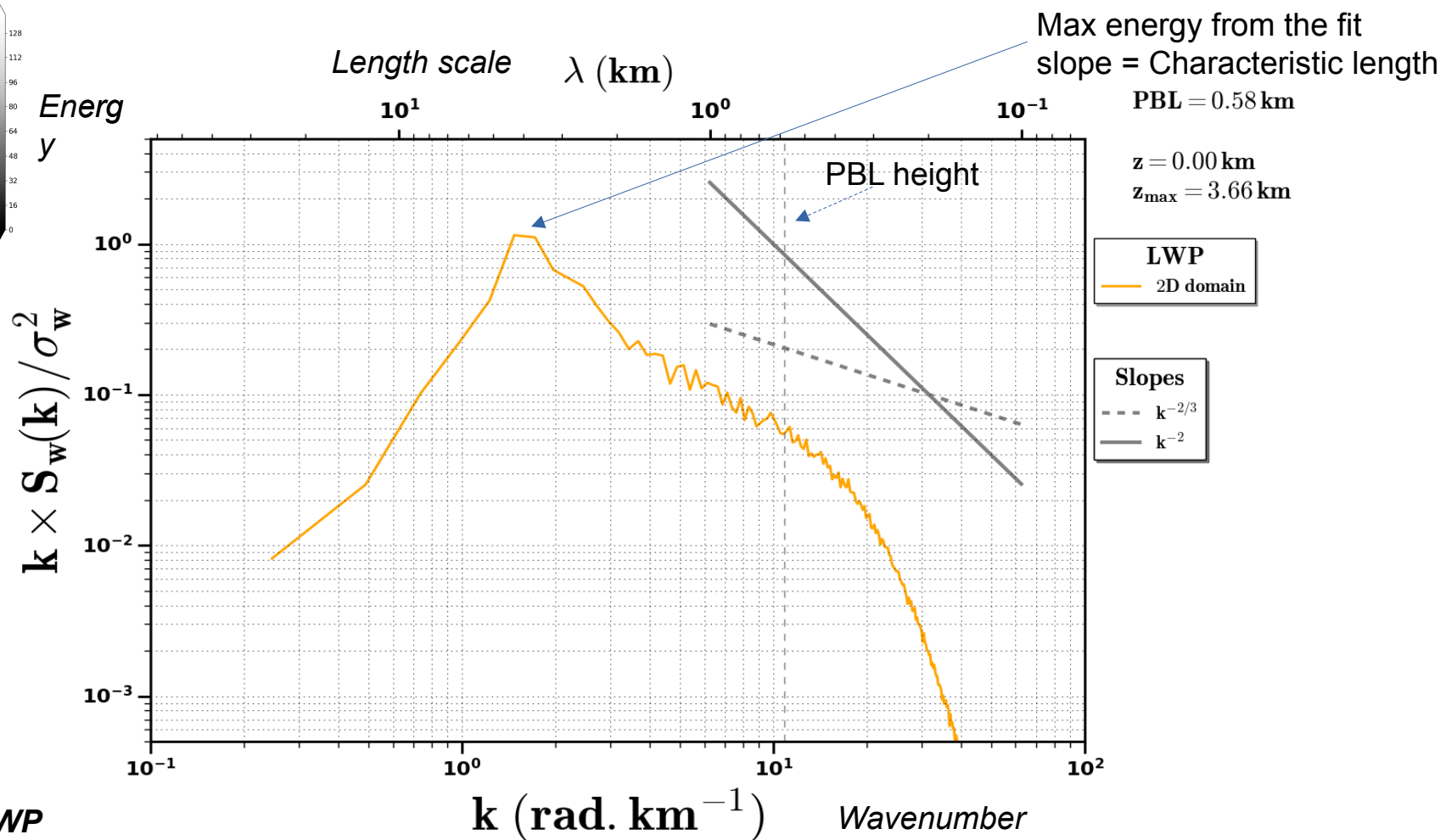
- Q1: Why have the stratocumulus a so large aspect ratio?
- Q2 :Is there some unified theory for **downdrafts' triggering** in all well-mixed layers?
- Q3: How should we **represent** downdrafts in climate models?
- Q4: Can we identify robust low-cloud feedback mechanisms?

*Morphology Of stratocumulus, BoundarY-layer DYnamics, and Climate Change (MOBYDYC) – ANR Project (2023-2027)*

# Cloud morphology (Large-eddy simulations)



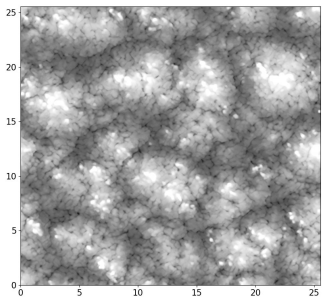
t+6h



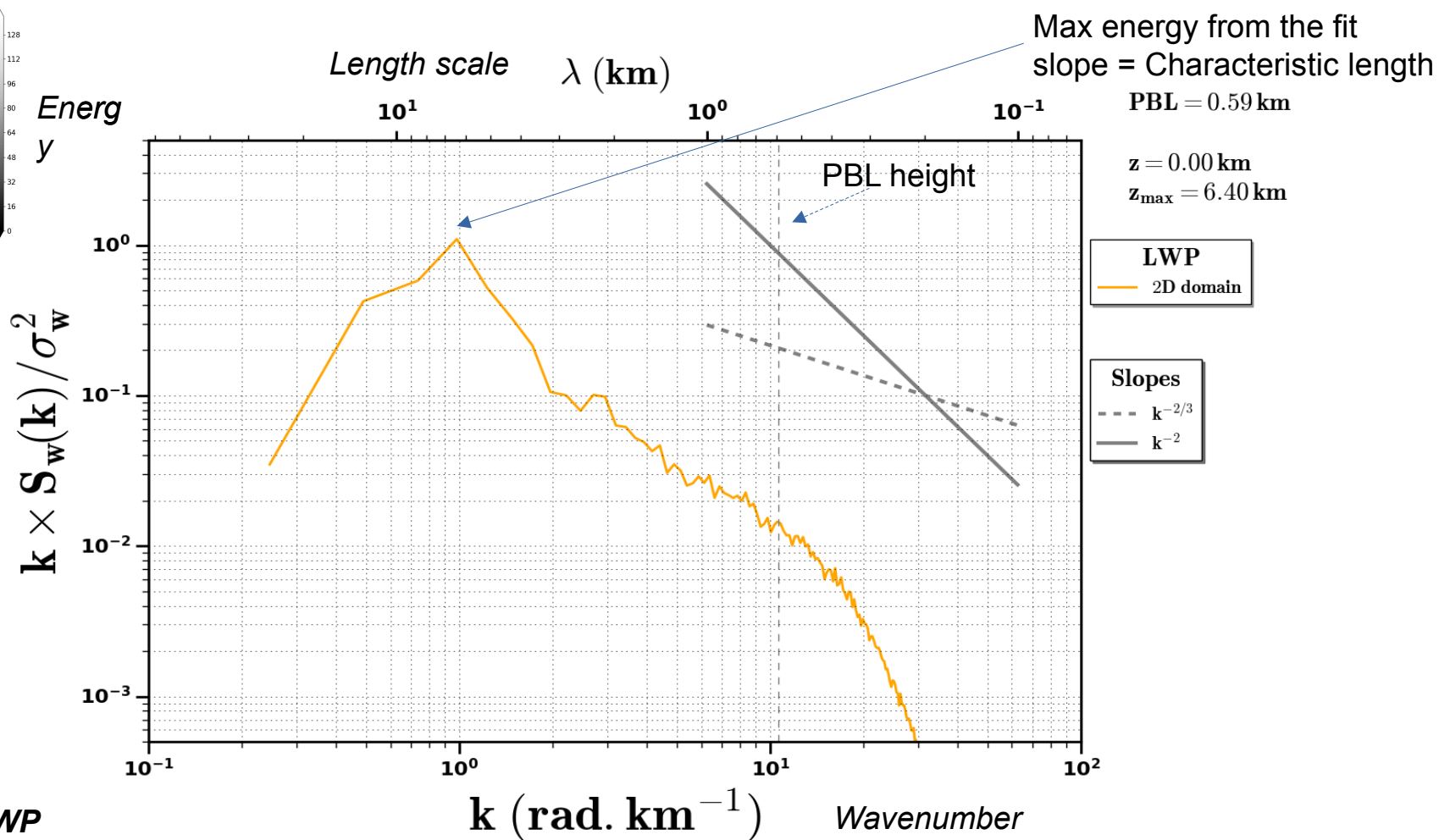
Power spectra of LWP

Cloudmetrics package by Janssens, Denby

# Cloud morphology (Large-eddy simulations)



t+12h

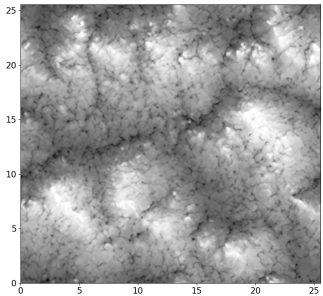


Power spectra of LWP

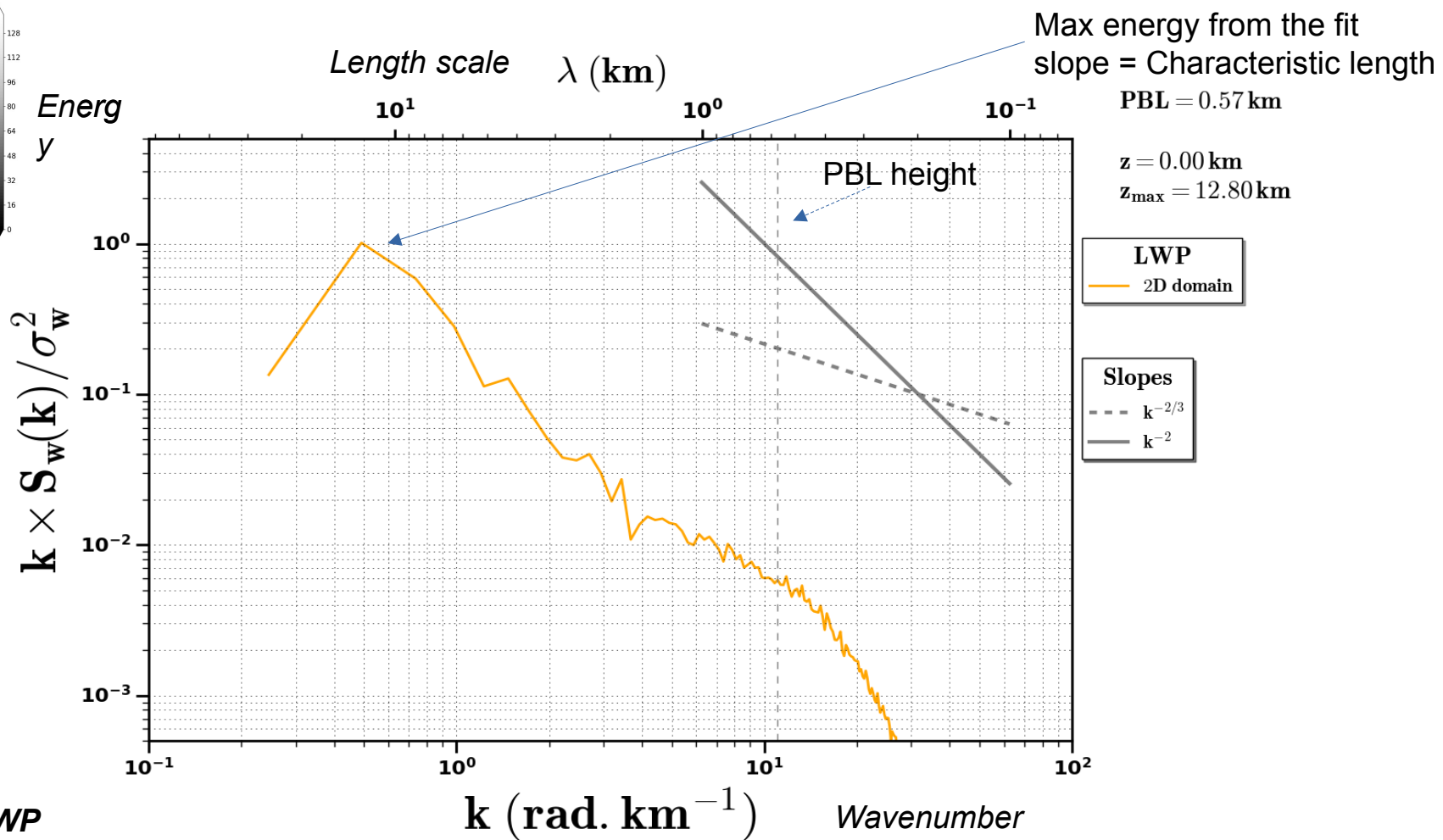
Cloudmetrics package by Janssens, Denby



# Cloud morphology (Large-eddy simulations)



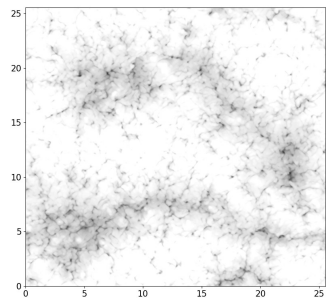
t+18h



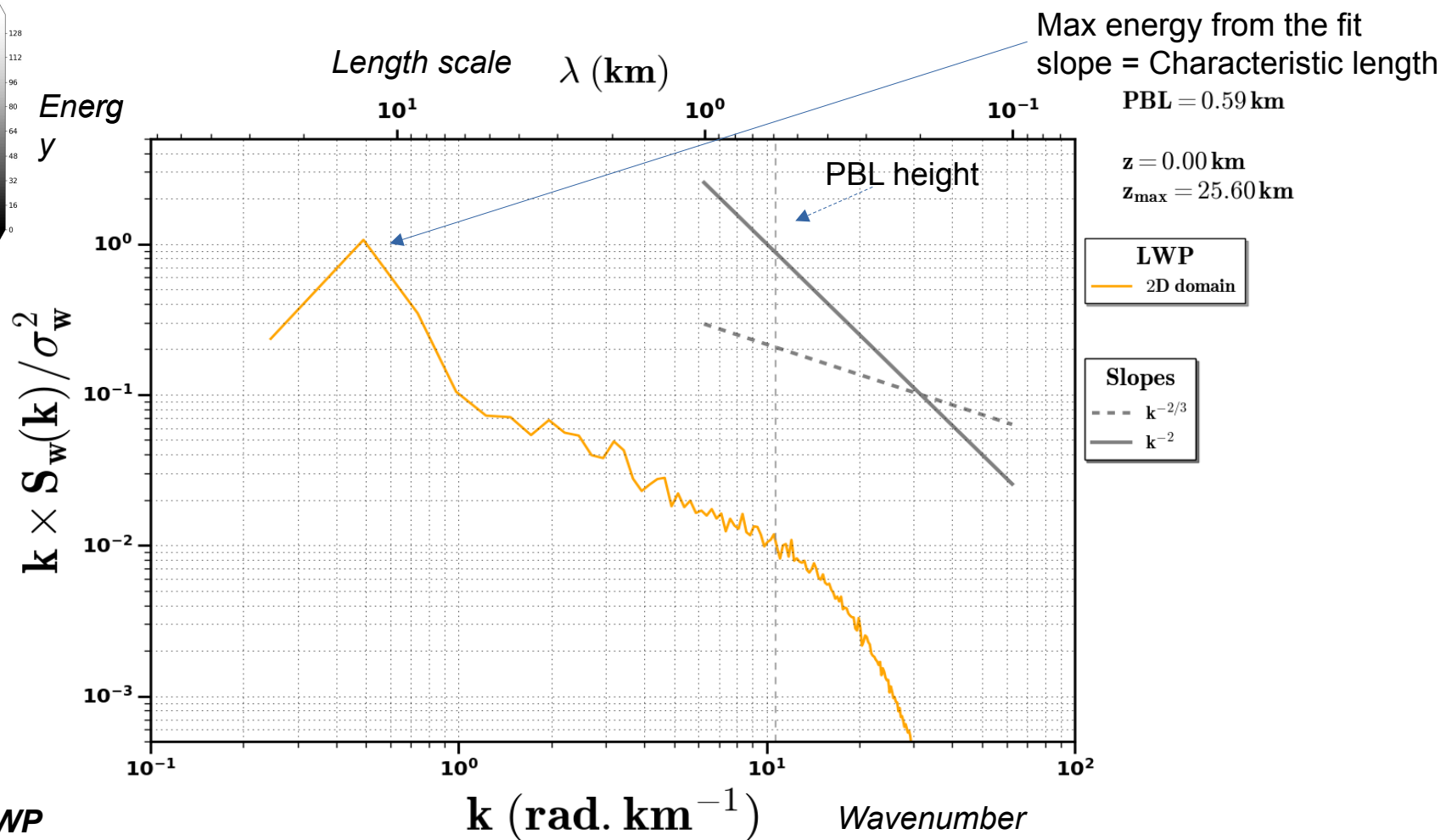
Power spectra of LWP

Cloudmetrics package by Janssens, Denby

# Cloud morphology (Large-eddy simulations)



t+24h



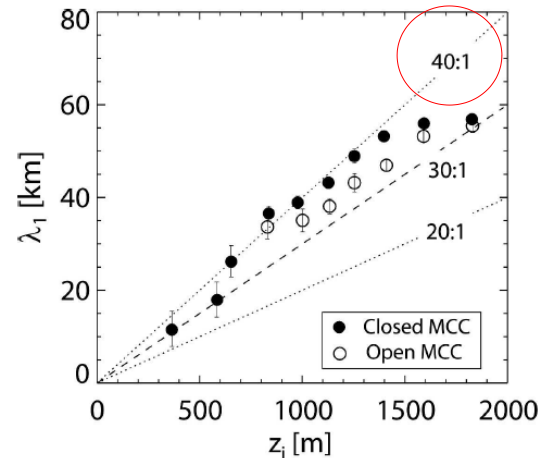
Power spectra of LWP

Cloudmetrics package by Janssens, Denby

# Cloud morphology (Large-eddy simulations)

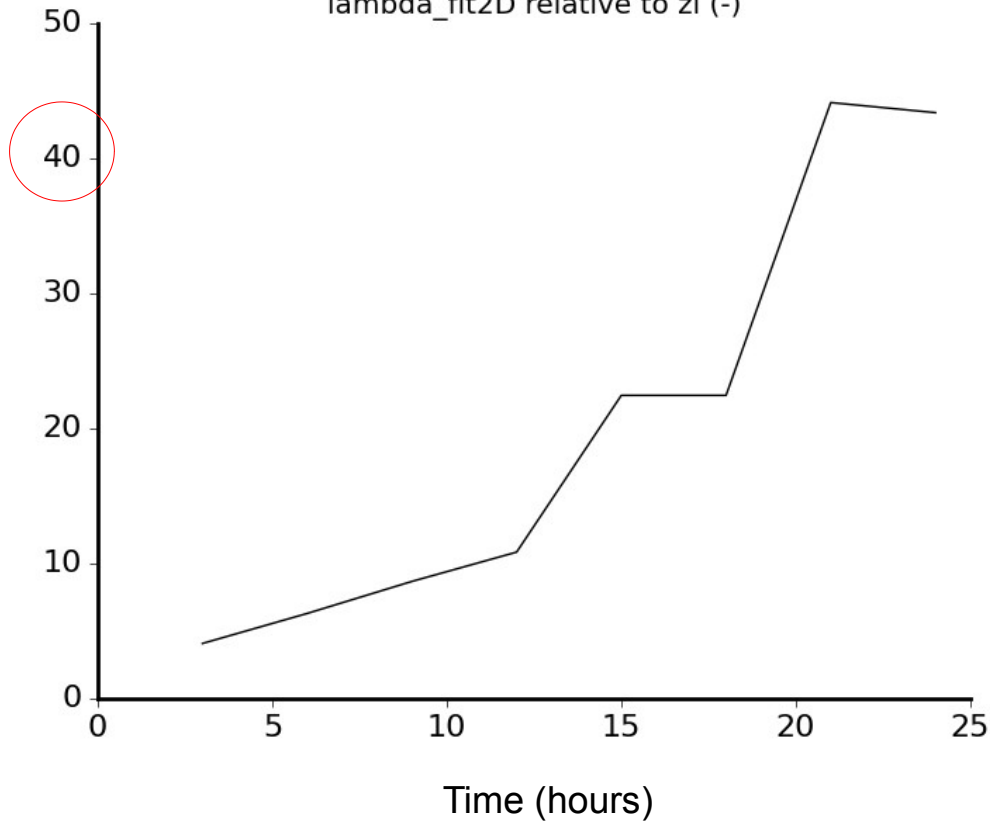
**Aspect ratio** (cell size / PBL height) of  
the Liquid Water Path field

Satellites  
Observations



Wood and Hartmann (2006)

lambda\_fit2D relative to zi (-)



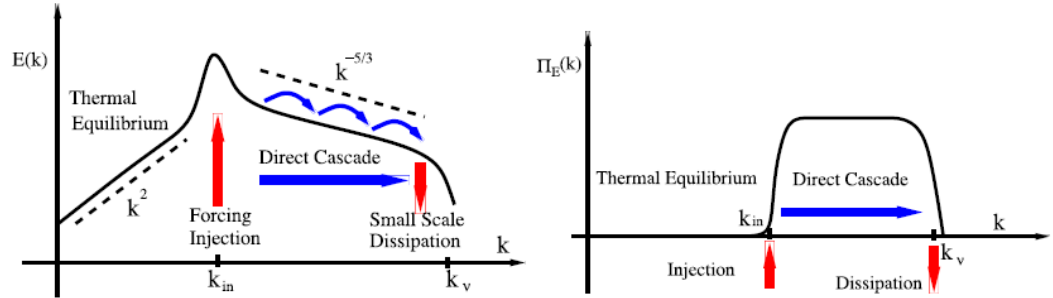
Upscale growth  
scale of  
structures

Why?

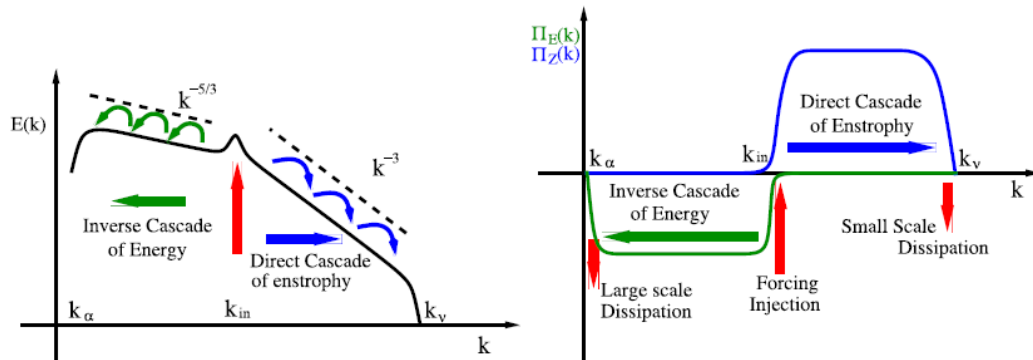
# Cloud morphology (theory)

A. Alexakis, L. Biferale / Physics Reports 767–769 (2018) 1–101

Classic 3D isotropic  
Kolmogorov cascade  
All vortices lose energy with  
surrounding smaller eddies



A. Alexakis, L. Biferale / Physics Reports 767–769 (2018) 1–101

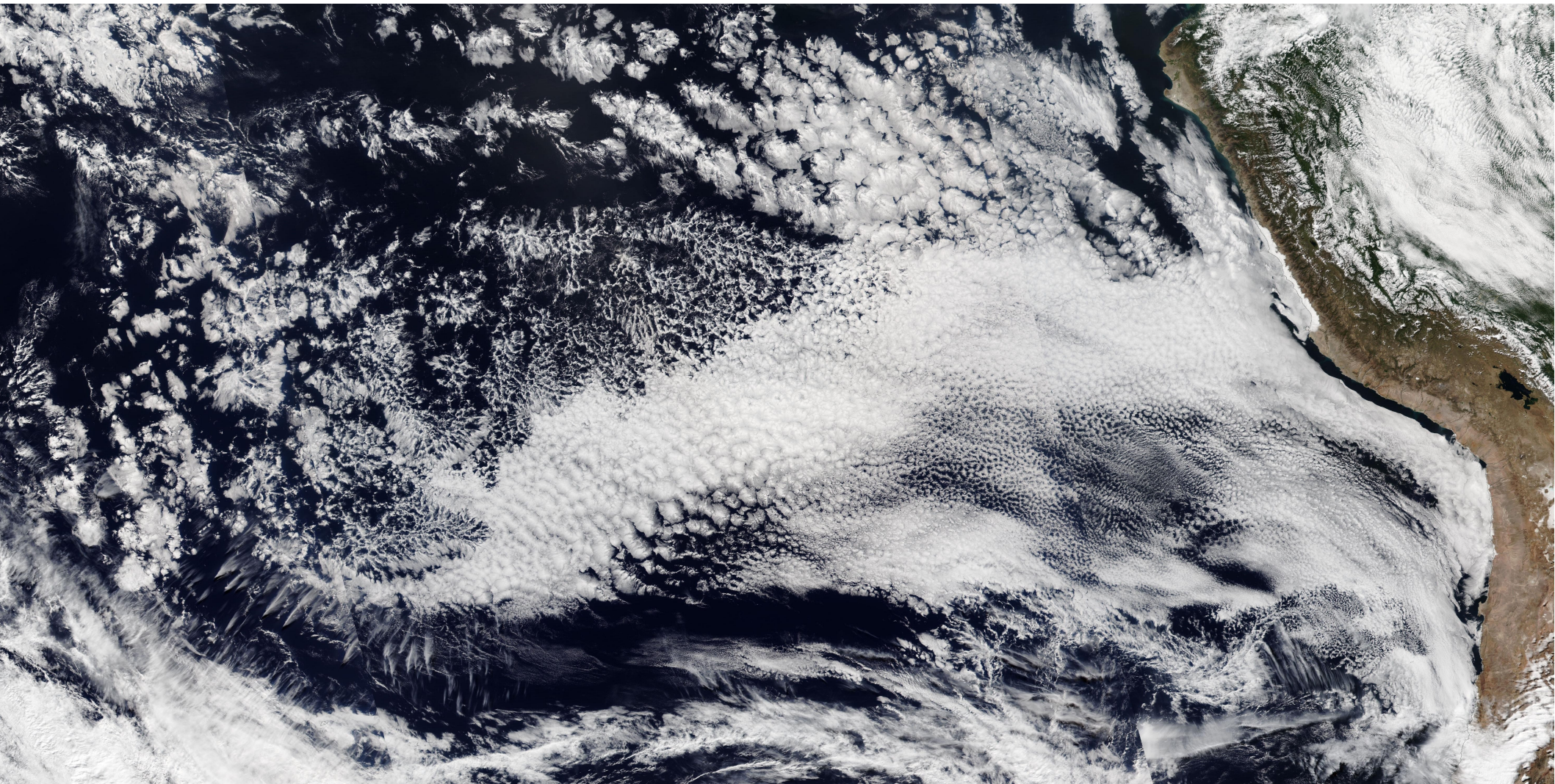


2D double cascade of energy.  
Inverse energy cascade  
suggest upscale growth above  
the length of energy injection

*Work in progress*



# Cloud morphology (observations)

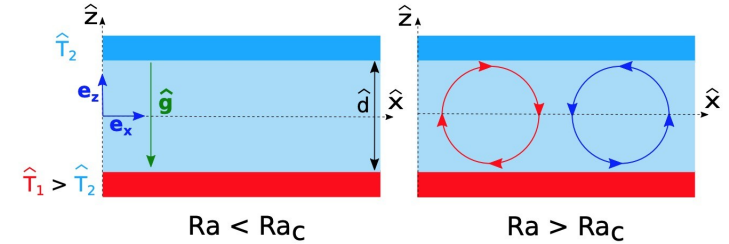




# An unified theory for atmospheric boundary layer organisation?

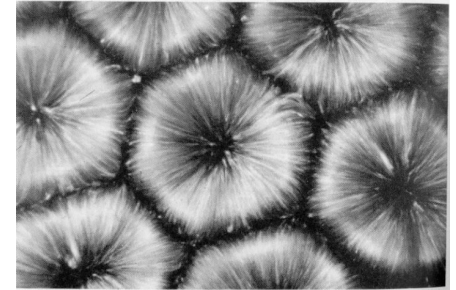
- Definition of the Rayleigh Bénard convection (RBC):

“A horizontal fluid layer of height  $d$  is confined between two thermally well conduction, parallels plates. When the difference  $DT = T_b - T_t$  between the bottomplate temperature and the top-plate temperature exceeds a critical value, the conductive motionless state is unstable and convection sets in. The simplest pattern which can occur is that of straight, parallel convection rolls” (Bodenschatz et. al, 10)



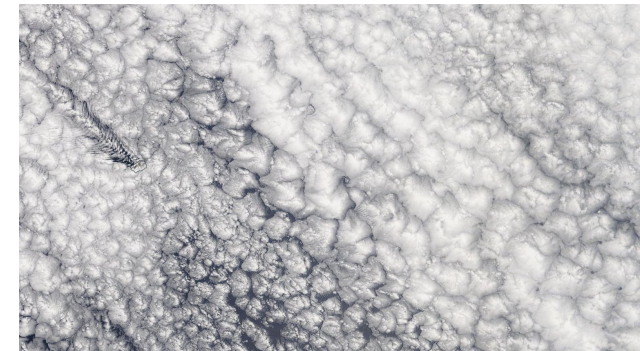
- **Similarities** between RBC and the Atmospheric Boundary Layer?

- ⌘ Fluid with high Rayleigh number (convection)
- ⌘ Warmer surface, colder troposphere (vertical T gradient)
- ⌘ Strong inversion as top plate?
- ⌘ Sensitivity of fluid properties to T and P solved by taking into account Non-Oberbeck-Boussinesq (NOB) effects (hexagons)



- **Differences** between RBC and the atmospheric BL?

- ⌘ The top-plate is **not rigid** (entrainment occur)
- ⌘ **Phase change** can modify RBC inside the convective layer and/or above (cumulus layer)
- ⌘ The **aspect ratio** of cells is larger than the RBC theory (30-50 for  $StCu \gg 1-2$ ).



# Conclusions

- Questions to go further
  - Q1: *Why has the stratocumulus a so large aspect ratio? Can we explain the upscale growth during the day? What is the exact role of decoupling in this evolution?*
    - **Power spectra** show an upscale growth of structures in clear-sky and stratocumulus
  - Q2 :*Is there some unified theory to understand downdrafts' triggering in all well-mixed layers?*
    - Structural organisation suggest that **Rayleigh-Bénard convection** is a good candidate
    - Still need to figure what are the exact role of entrainment, condensation, heterogeneities in modifying the canonical RBC
  - Q3: *How should we represent downdrafts in climate models?*
    - **Coherent subsiding structures** need to be represented, compensating subsidence not enough
  - Q4: *Can we highlight robust low-cloud feedback mechanisms?*
    - Not yet

## 2026 Workshop idea:

“Theoretical advances in understanding the organization of atmospheric (oceanic?) boundary layers” (or something like that) - Link with GDR Defis théoriques, DEPHY, GASS, Annual Workshop Organisation Convection



Thank You

